Use of Satellite Data for Gridded SST Analyses of Pre-Satellite Period

Alexey Kaplan

Lamont-Doherty Earth Observatory of Columbia University, Palisades, NY 10964, USA

Alicia Karspeck

National Center for Atmospheric Research, Boulder, CO 80307, USA

OUTLINE

Historical data availability

- Optimal Interpolation (OI) as gridding and analysis method
- Estimation of input covariances (for observational and background error)
- Utility of the space reduction for coarse gridding of historical in situ data
- Importance of satellite data for estimating parameters of higher resolution OI
- Representing uncertainty by samples from the posterior distribution





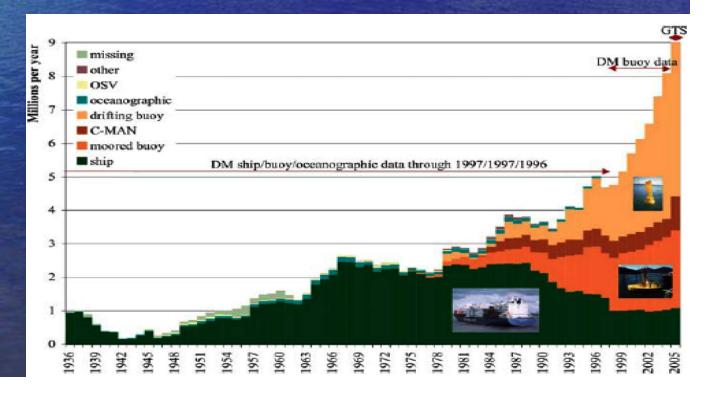
From Woodruff et al. [2008], In Climate Variability and Extremes during the Past 100 Years, Bronniman et al. (eds.)

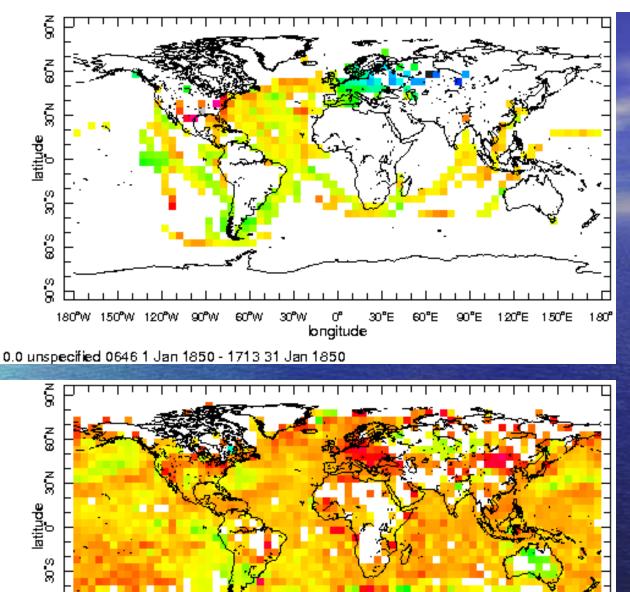
Transition to the modern Ocean Observing System











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brigitude

30'

30°E

€0°E

90°E

120°E 150°E

180°

Jan 1850

In situ data: HadSST2 [Rayner et al., 2006]

Jul 2007

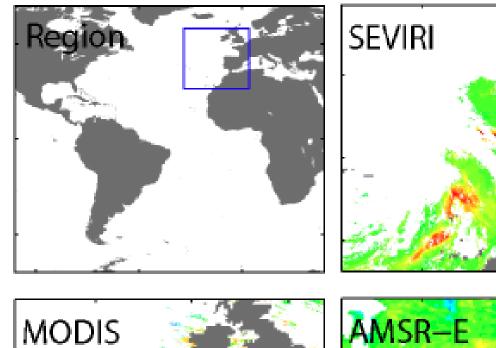
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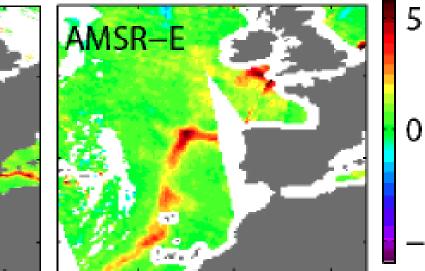
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> > 180°W

Satellite Observations





ΔT (K)

Donlon et al. [2010], OceanObs'09, Community White Paper

PROBLEM

Richness and volume of these data sets notwithstanding, as descriptors of detailed historical SST variability they are quite incomplete, affected by large errors, or else of rather short time coverage.

Yet many applications require SST fields interpolated onto a regular grid, with no spatial or temporal gaps, and uncertainty estimates.

Optimal Interpolation (OI)

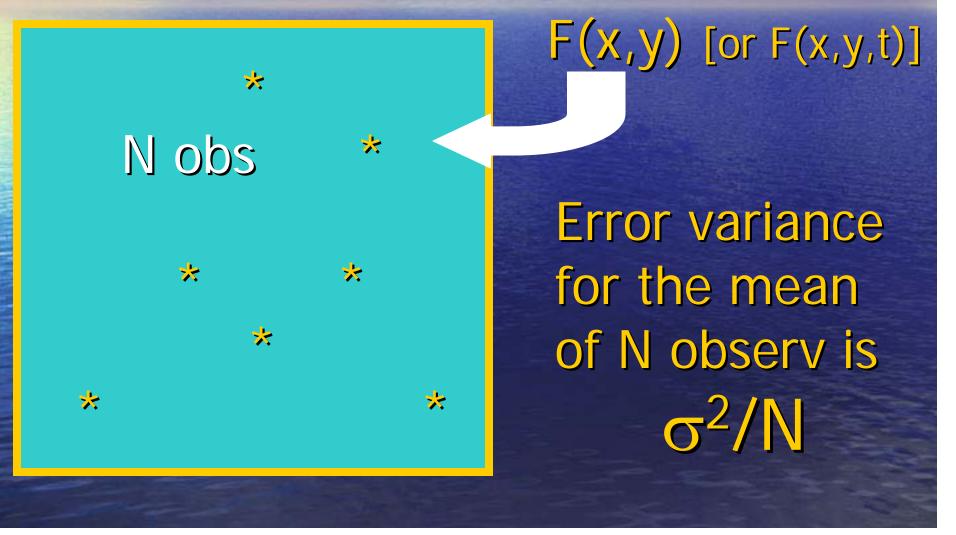
 $T = T_{B} + e_{B}$ $HT = T_{o} + e_{o}$ $< e_{B} > = < e_{o} > = < e_{B}e_{o}^{T} > = 0$ $< e_{B}e_{B}^{T} > = C \leftarrow bckgr \ err \ (signal) \ cov$ $< e_{o}e_{o}^{T} > = R \leftarrow observ \ err$

covariance

Solution minimizes the cost function $S[T] = (HT - T_{o})^{T}R^{-1}(HT - T_{o}) + (T - T_{B})^{T}C^{-1}(T - T_{B})$ $T_{OI} = P_{OI}(H^{T}R^{-1}T_{o} + C^{-1}T_{B}), P_{OI} = (H^{T}R^{-1}H + C^{-1})^{-1}$

What we infer from the data, given the Ol assumptions $T \sim \mathcal{N}(T_{OI}, P_{OI}),$ where $T_{OI} = P_{OI}(H^{T}R^{-1}T_{O} + C^{-1}T_{B})$ $= T_{\rm R} + CH^{\rm T}(HCH^{\rm T} + R)^{-1}(HT_{\rm O} - T_{\rm R}),$ $P_{OI} = (H^{T}R^{-1}H + C^{-1})^{-1}$ $= C-CH^{T}(HCH^{T}+R)^{-1}HC$ Matrix dimensions are R: Nobs XNobs , C: Ng XNg, i.e., for 5° grid $N_a \sim 2000$, for 1° grid $N_a \sim 50,000$ For sparsely sampled historical in situ data, a sparse grid (e.g. 5° grid size) makes sense. Sparse grids best represent largest scales of variability. Hence the Reduced Space (RS) approach:

T = T' + T'' = C = C' + C'',where T' is a linear combination of a few large scale patterns and C' is of low rank. If T' and T'' are assumed independent, there is a cheap way to compute T' and its error covariance P' corresponding to the OI solution cheaply (RSOI) For a sparse grid observational error covariance R is usually assumed diagonal; its element are estimated as uncertainties in the grid box averages

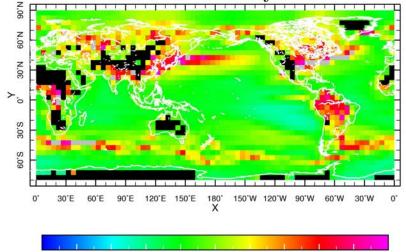


Finer data coverage provided by satellite data is helpful for better estimates of R.

(right) modeling sampling error in grid box averages of in situ data

(bottom) verification

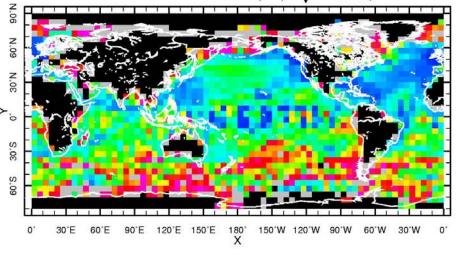
Single observation SST sampling+measurement error, ${}^{o}C$, inside $5^{o} \times 5^{o}$ monthly bins

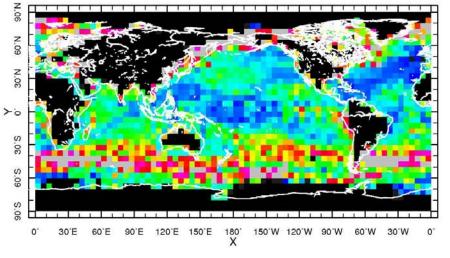


 $\begin{array}{ccc} \mbox{Modeling in situ data error for 5}^o \mbox{ bins} \\ \mbox{Modeled as } \langle \sigma / \sqrt{n_{\rm obs}} \rangle & \mbox{Actual MODIS-ICOADS STD} \end{array}$

08

1.2





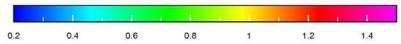
2.4

2.8

3.2

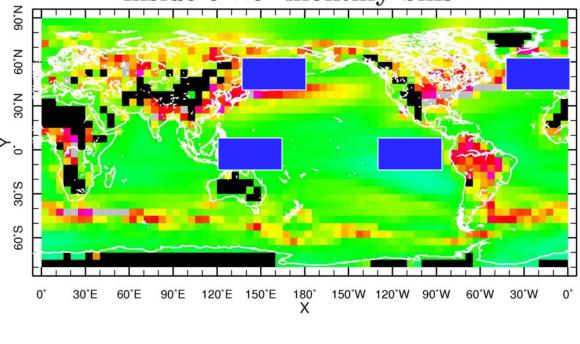
3.6

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0	0.1	0.2	0.3	0.4	0.5	0.6	0.7	0.8	0.9	1



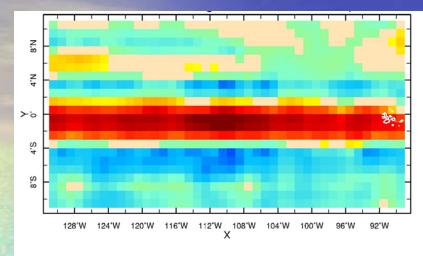
Finer data coverage is helpful for better estimates of R. If we would like to increase the resolution, a lot of detail is needed for estimating R and C"

Single observation SST sampling+measurement error, ${}^{o}C$, inside $5^{o} \times 5^{o}$ monthly bins





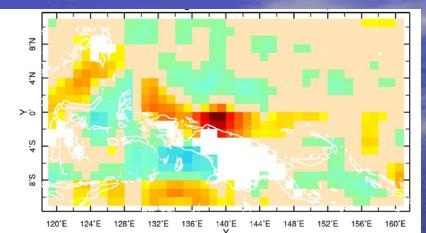
Truncation error autocovariance patterns: Eastern Equatorial Pacific at 110W Western Equatorial Pacific at 140E



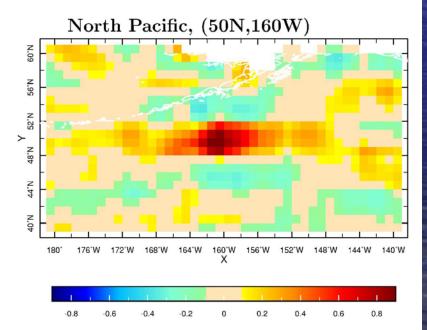
1	1.1.1							
-0.8	-0.6	-0.4	-0.2	0	0.2	0.4	0.6	0.8

North Atlantic, (50N, 20W) 0 N z 44°N 40'N 40'W 20'W 4'W 0" 36°W 32°W 28'W 24'W 16'W 12'W 8'W Х

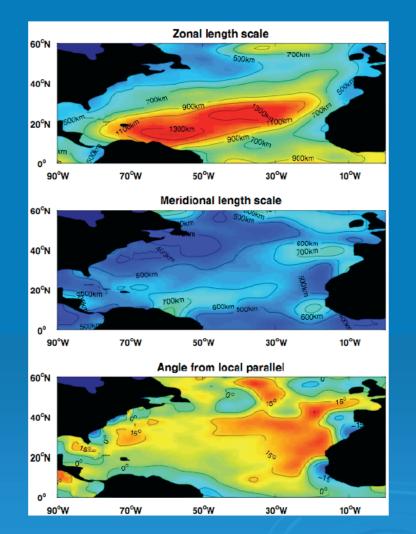


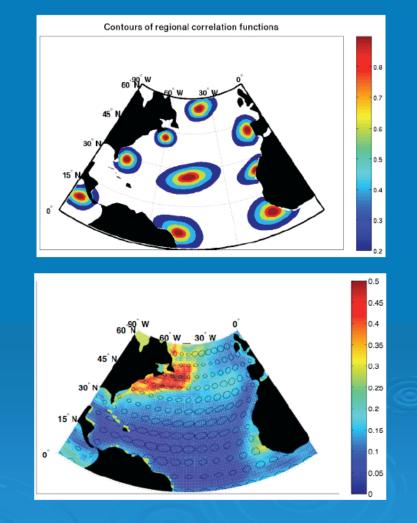


				. n			1.1	1 1
-0.8	-0.6	-0.4	-0.2	0	0.2	0.4	0.6	0.8



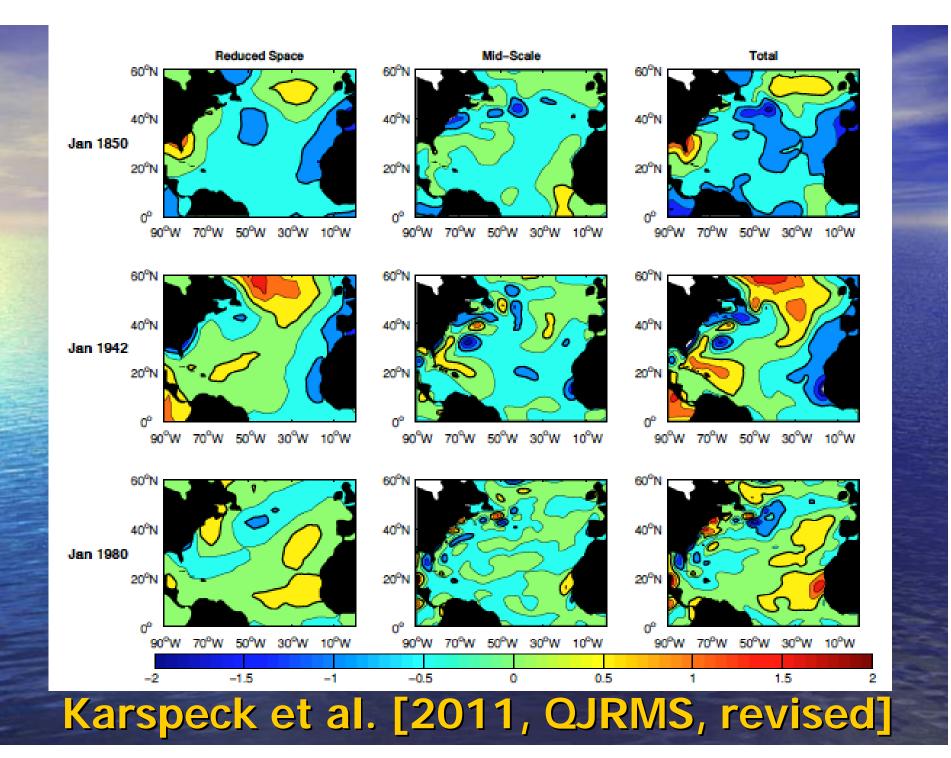
Covariance structures in the mid-scale SST

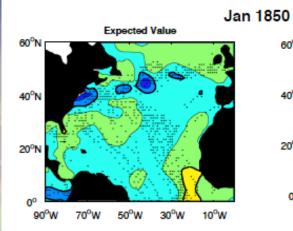


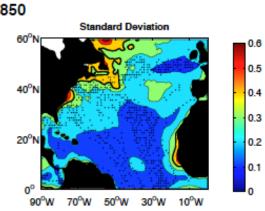


Parameters estimated from NCEP OI from 1981-present via maximum likelihood

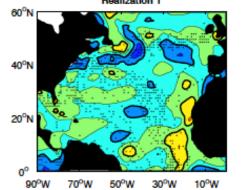
Karspeck et al. [2011, QJRMS, revised]

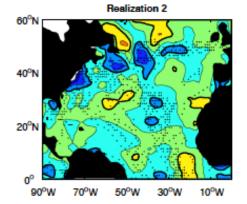




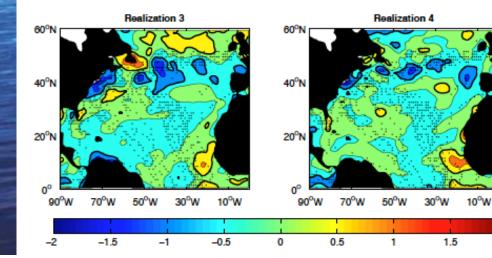




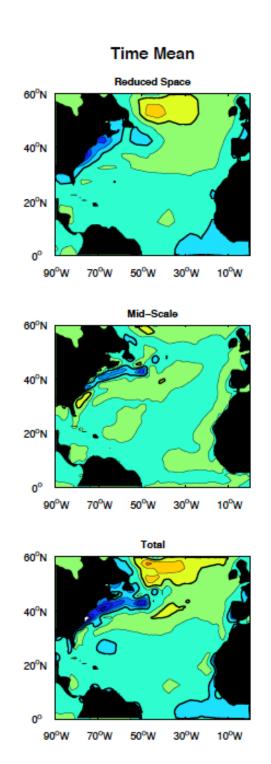




2



Karspeck et al. [2011, QJRMS, revised]



1

0.8

0.6

0.4

0.2

0

-0.2

-0.4

-0.6

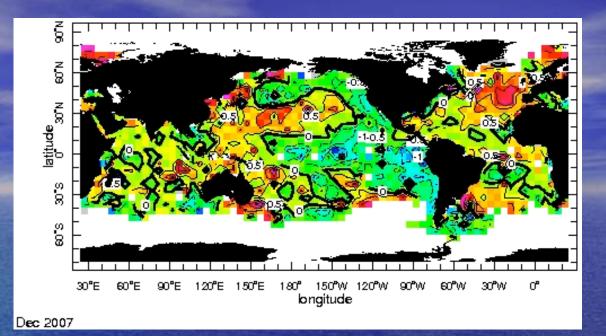
-0.8

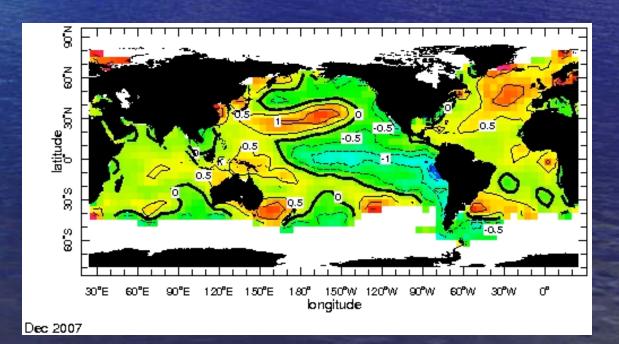
-1

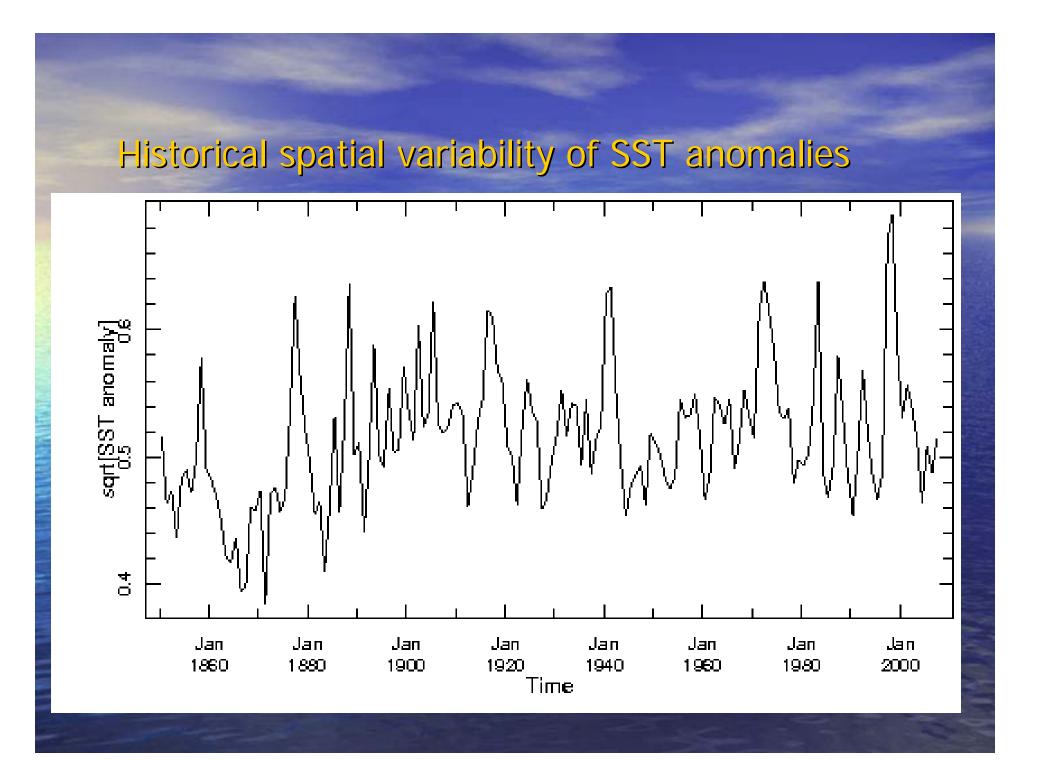
10⁰ 10 10 10 5 10 15 20 25 30 35 40 45 50 Karspeck et al. [2011, QJRMS, revised] $LT \sim N(LT_{OI}, LP_{OI}L^{T})$

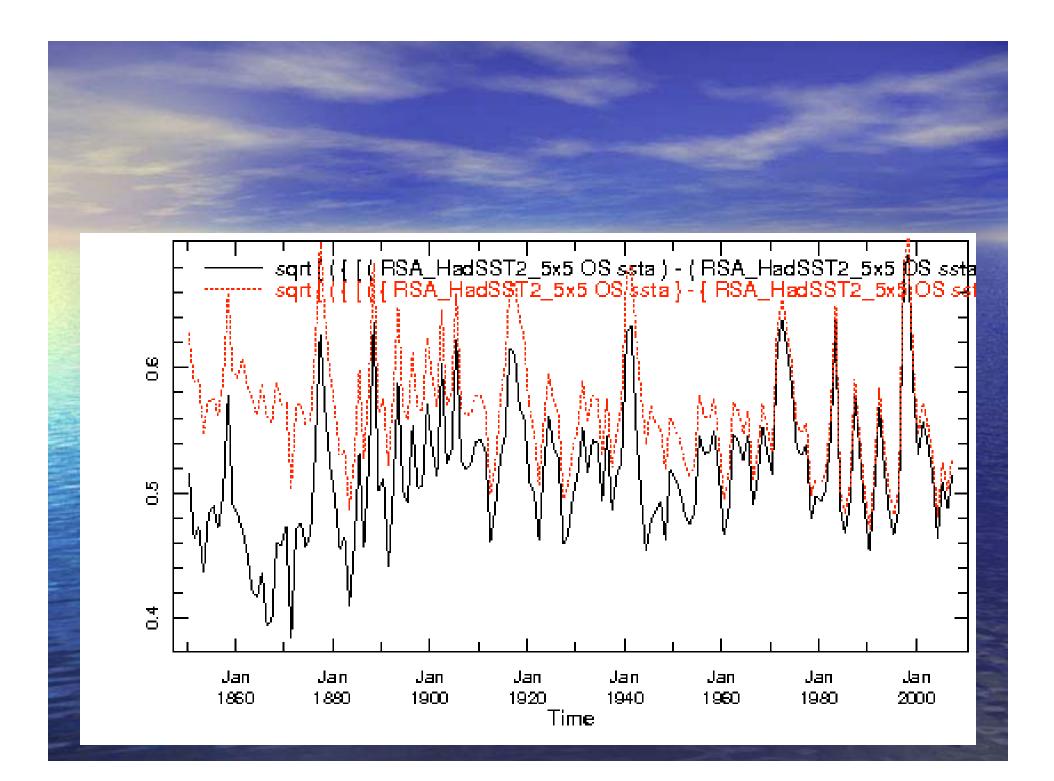
 $\langle T^{T}ST \rangle = T_{OI}^{T}ST_{OI} + Tr(SP_{OI})$

Reduced Space Optimal Smoother applied to HadSST2 data set of the U.K. MetOffice









Conclusions:

 High-resolution data satellite data can be used for estimating data covariance at different scales and then be used in efficient multi-scale optimal analysis procedures.
Data covariance is spatially non-stationary at both large and small scales and can be

modeled as such.

3. Representation of the uncertainty by a set of samples from the posterior distribution is viable and useful

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