



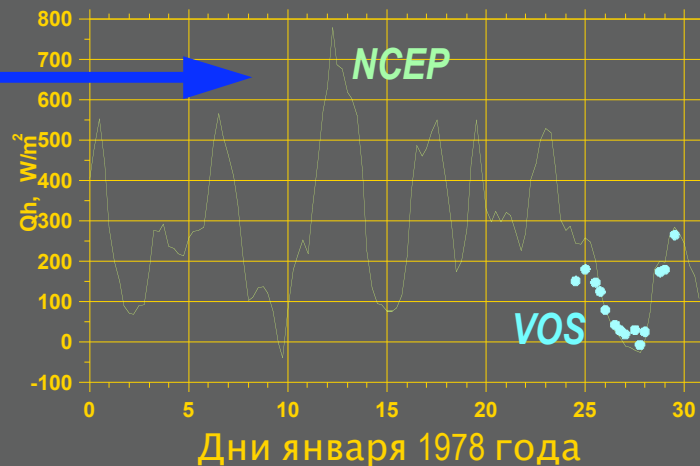
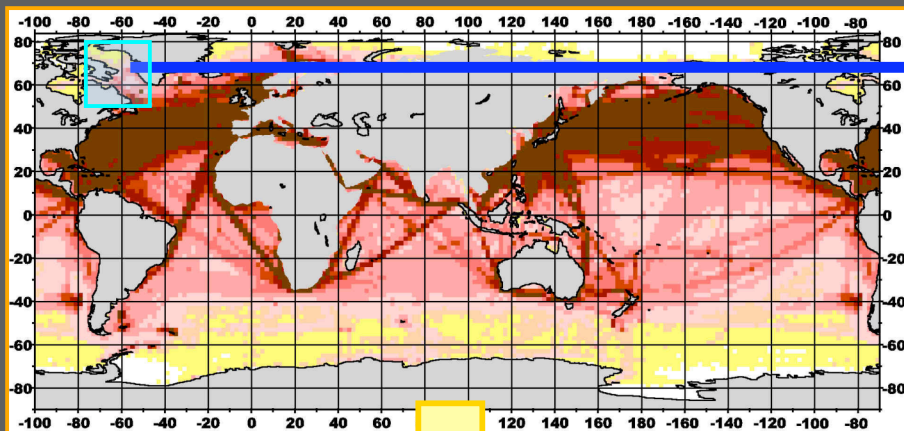
Reconstruction of interdecadal variability of air-sea interaction in the Atlantic 1880-2004: links to atmospheric circulation patterns

Sergey Gulev, Konstantin Belyaev, Thomas Jung, Eberhard Ruprecht

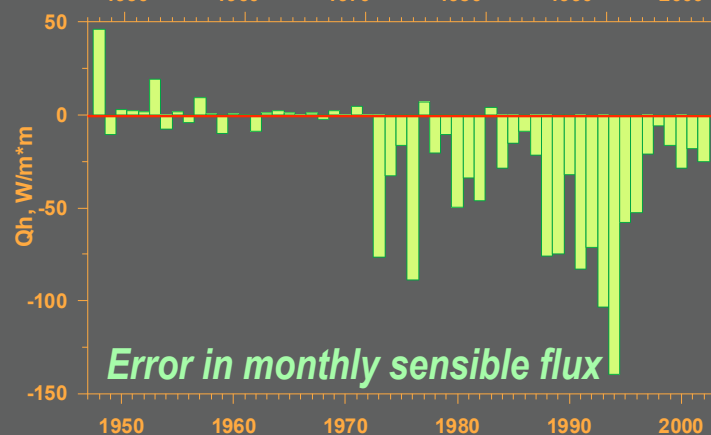
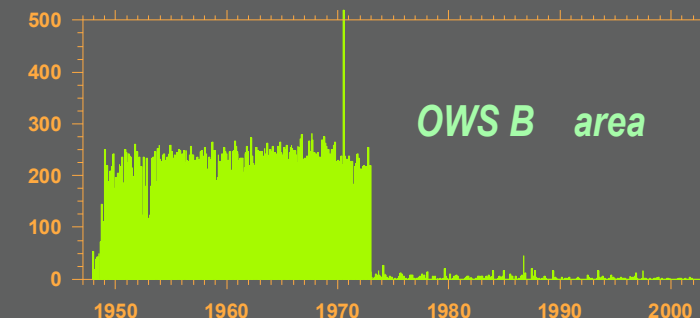
Outline:

- ❑ Introduction: Uncertainties of surface flux fields : importance of sampling errors
- ❑ A way to account for sampling uncertainty and to minimize it: using modified FT-distribution (2eWPDF)
- ❑ Homogenized 100-yr flux time series in the North Atlantic from VOS (ICOADS)
- ❑ Surface heat fluxes and the NA variability patterns
- ❑ Conclusions

The nature of sampling bias in air-sea fluxes

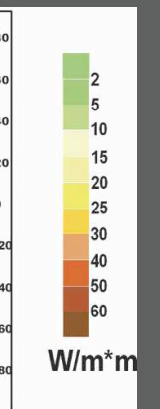
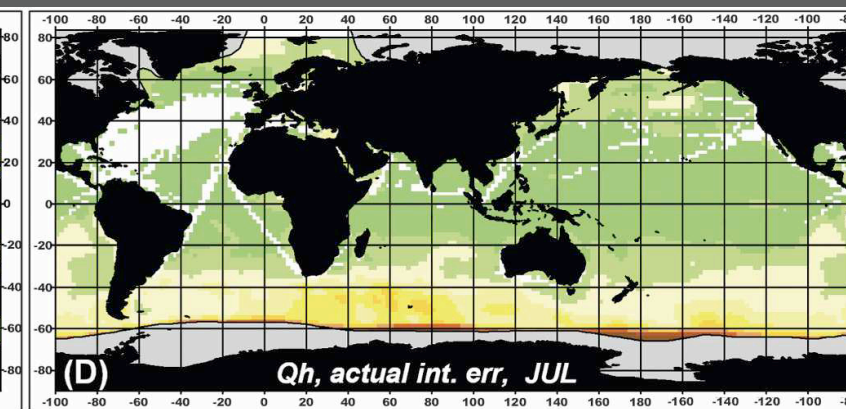
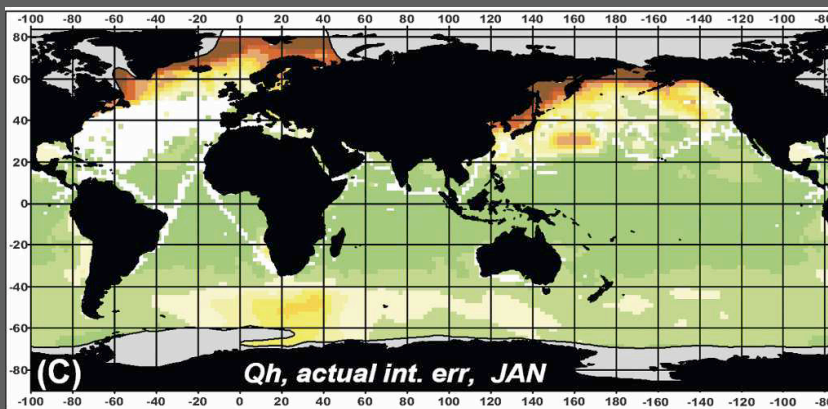
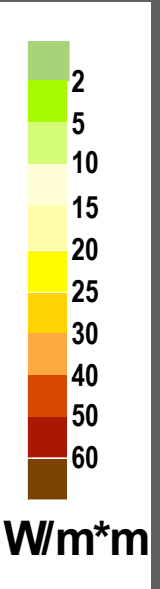
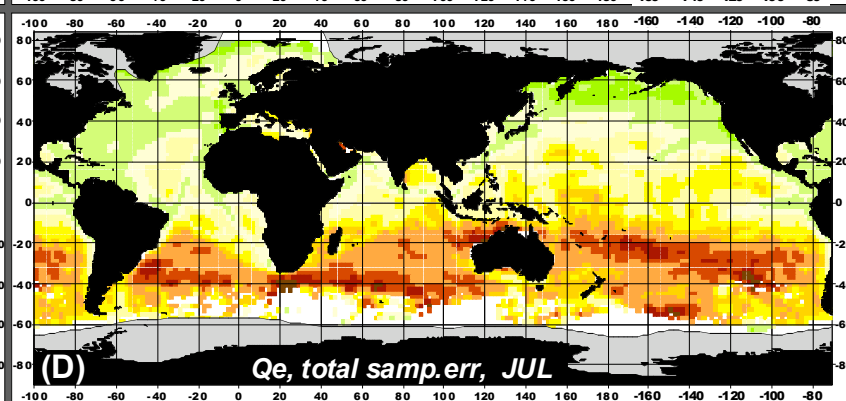
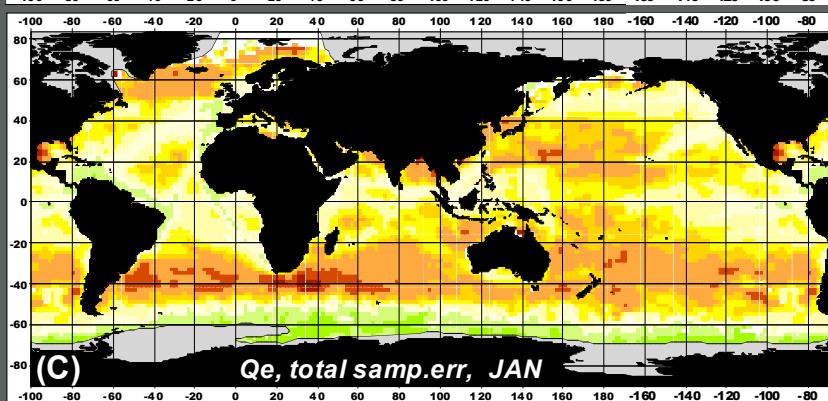
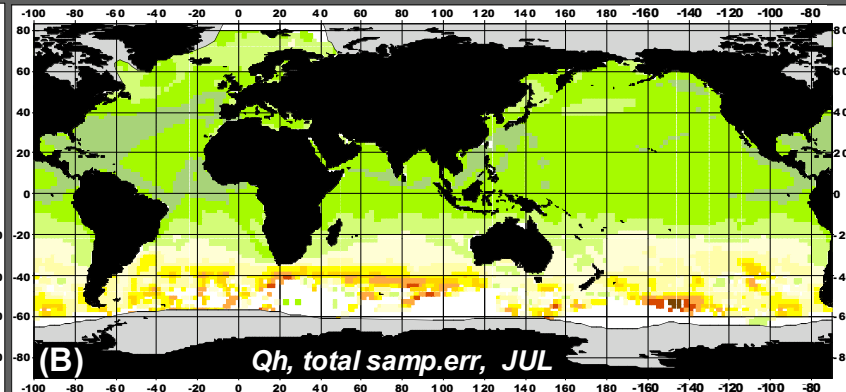
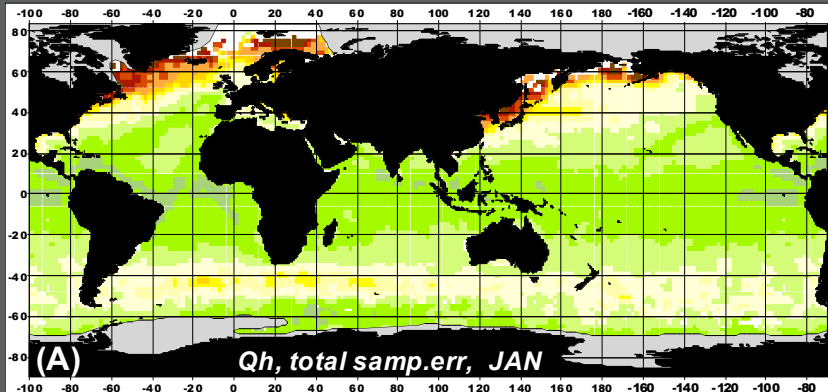


“VOS-like” re-sampling
of NWP in order to
quantify sampling bias

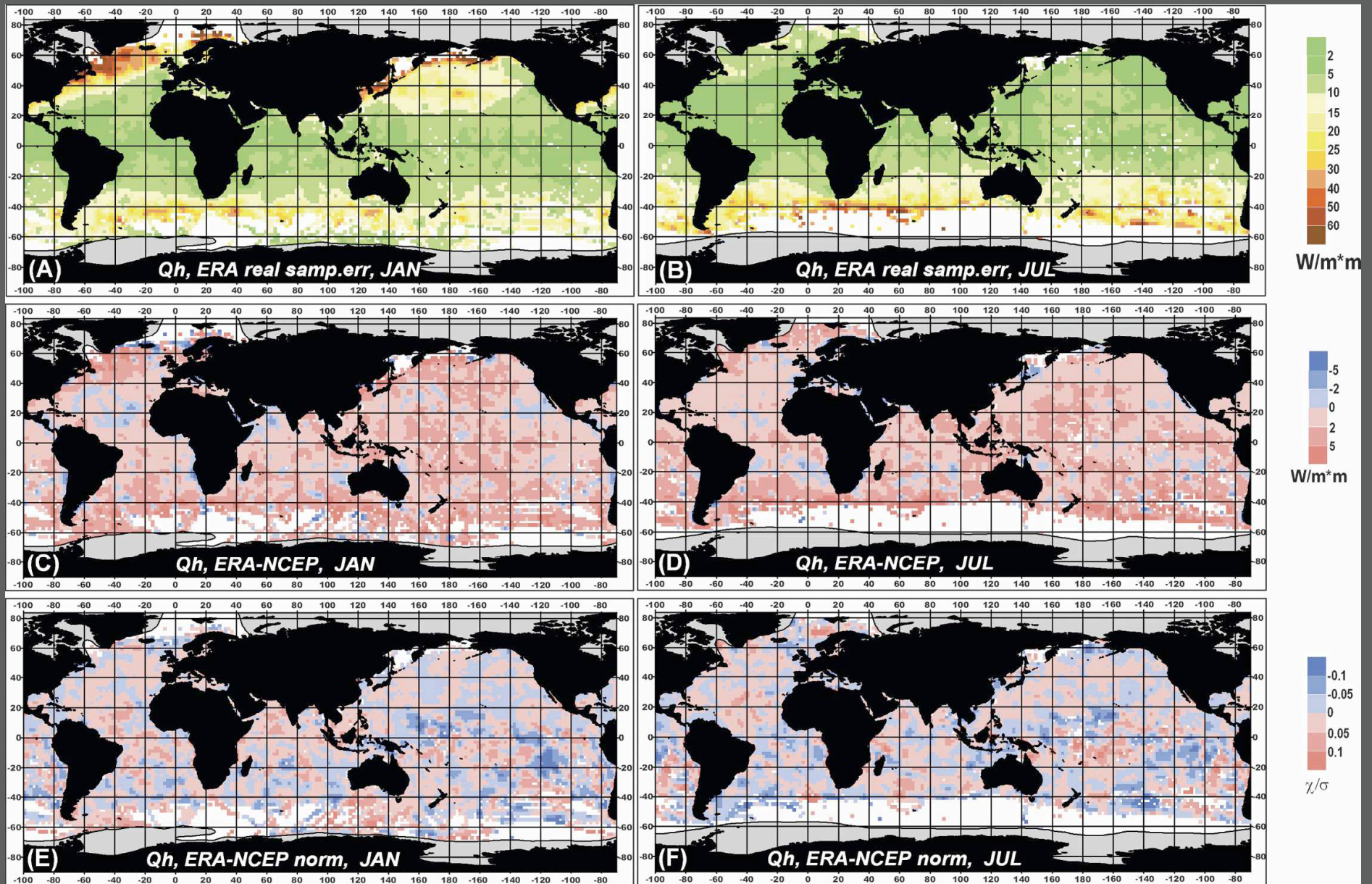


Gulev et al., 2007a,b *J. Climate*

Magnitude of sampling error

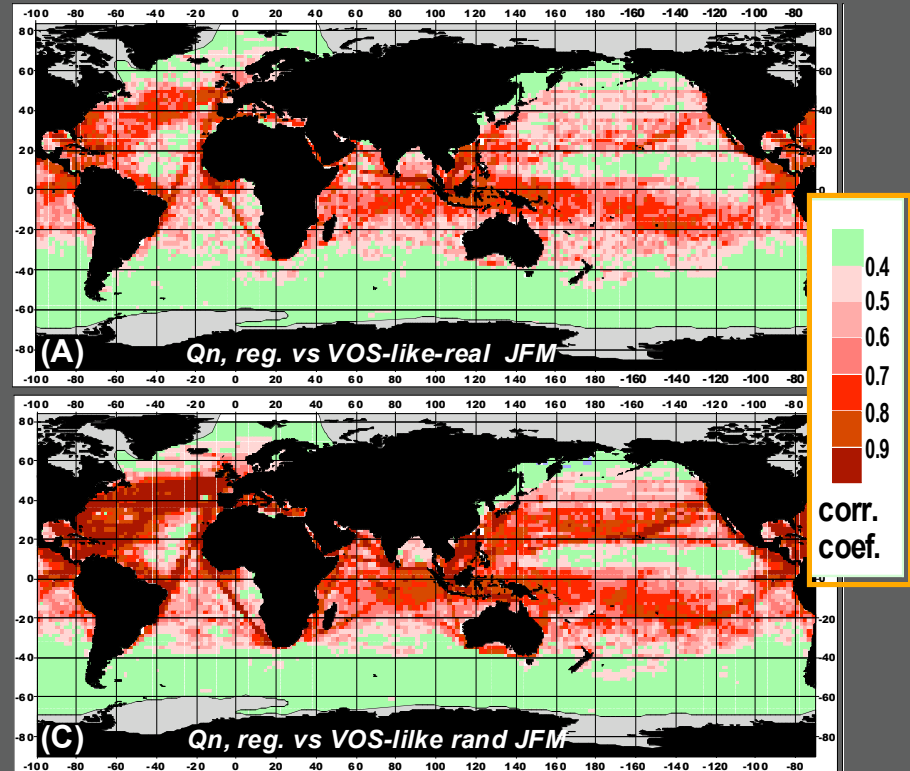
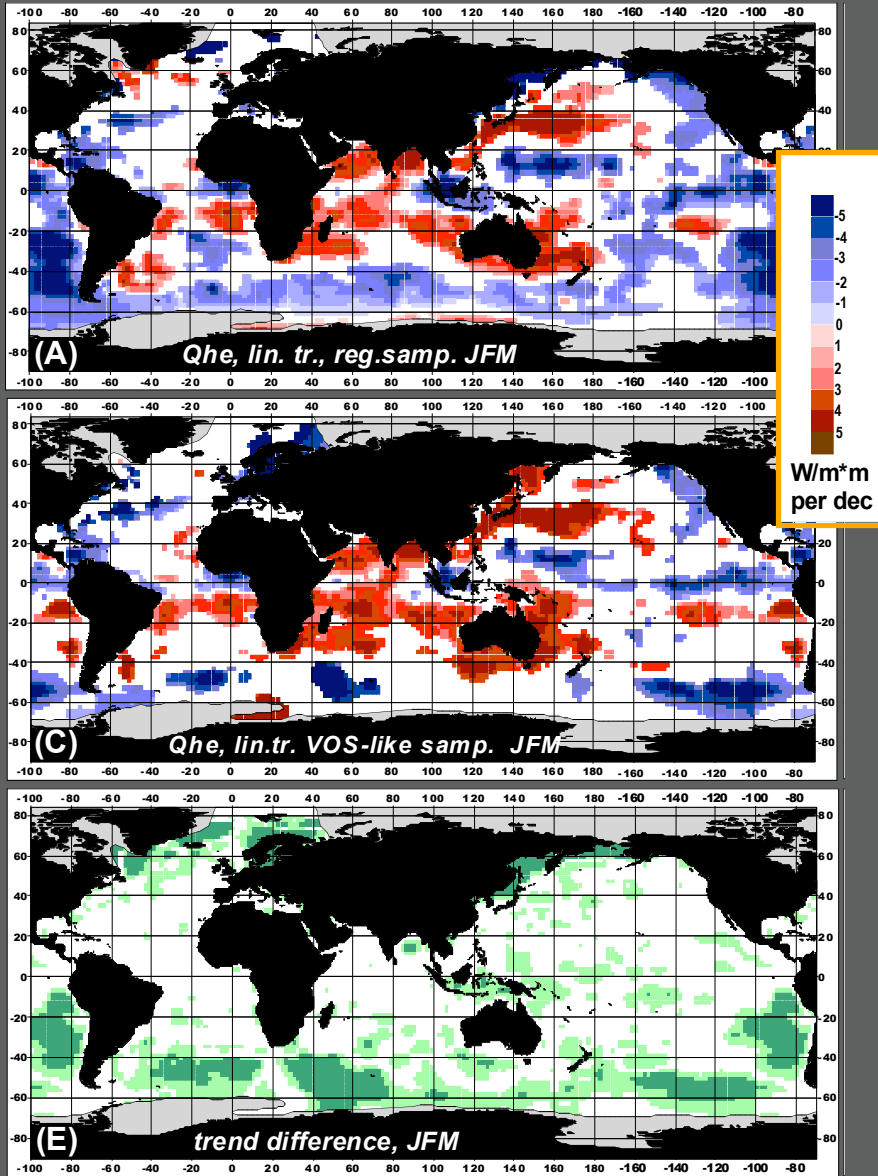


Sampling errors from different NWP



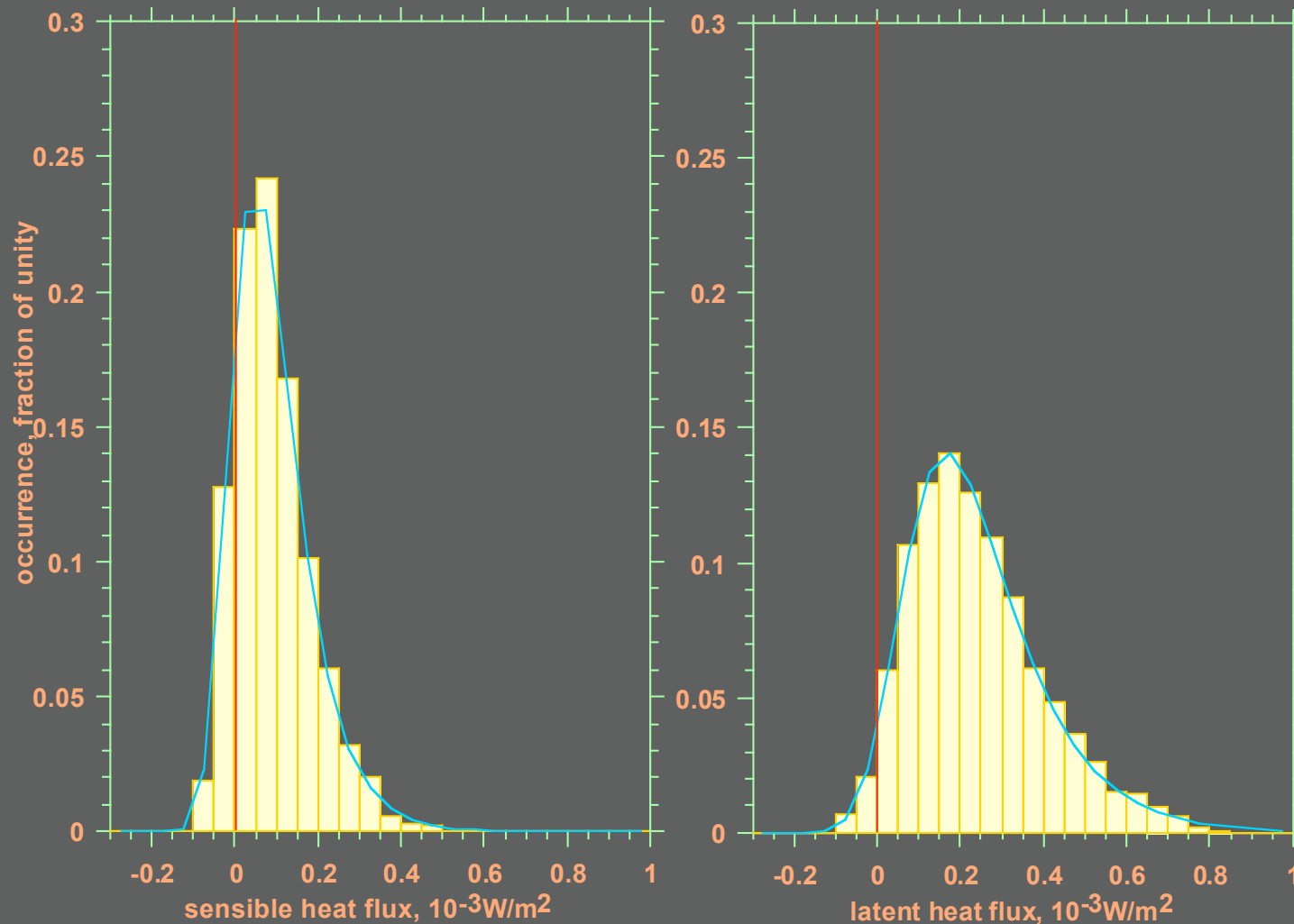
Sampling affecting trends and interannual variability in surface fluxes

- ❑ Changing trend sign in subpolar NA
- ❑ No consistency in interannual variability



How to overcome sampling error?

Although it is clear how to quantify sampling error, it is unclear how to minimize it → PDF is needed



Modified FT-distribution (2ePDF)

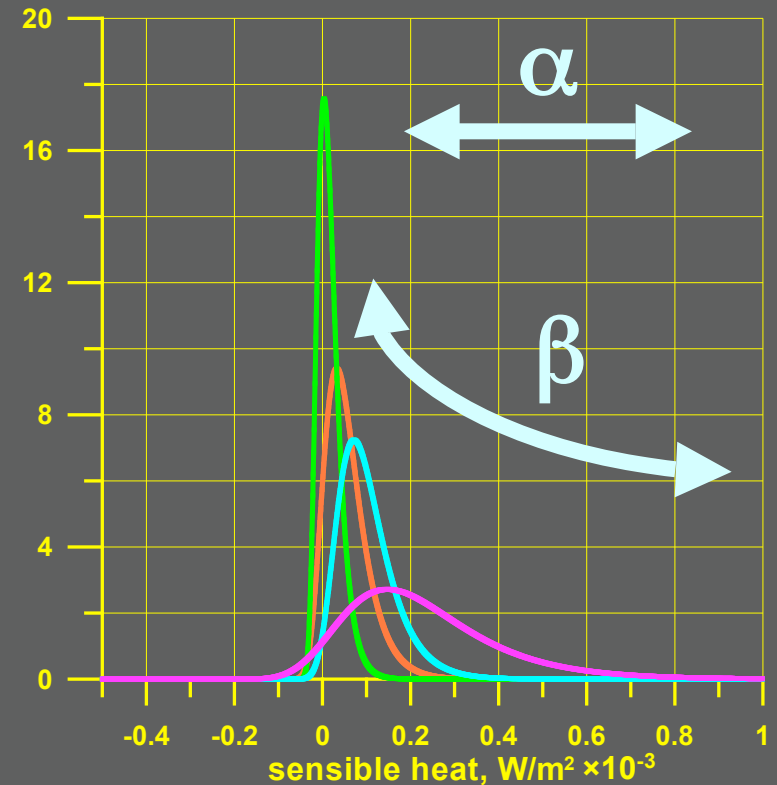
PDF: $P(x) = -(\alpha \cdot \beta) \cdot e^{\beta x} \cdot e^{-\alpha \cdot e^{\beta x}}$, $\alpha > 0, \beta < 0$

Mean, variance, modal value:

$$\bar{x} = \int_{-\infty}^{\infty} P(x) \cdot x dx = \frac{C + \ln \alpha}{-\beta}$$

$$\text{var } x = \int_{-\infty}^{\infty} P(x) \cdot x^2 dx - \bar{x}^2 = \frac{\pi^2}{6\beta^2}$$

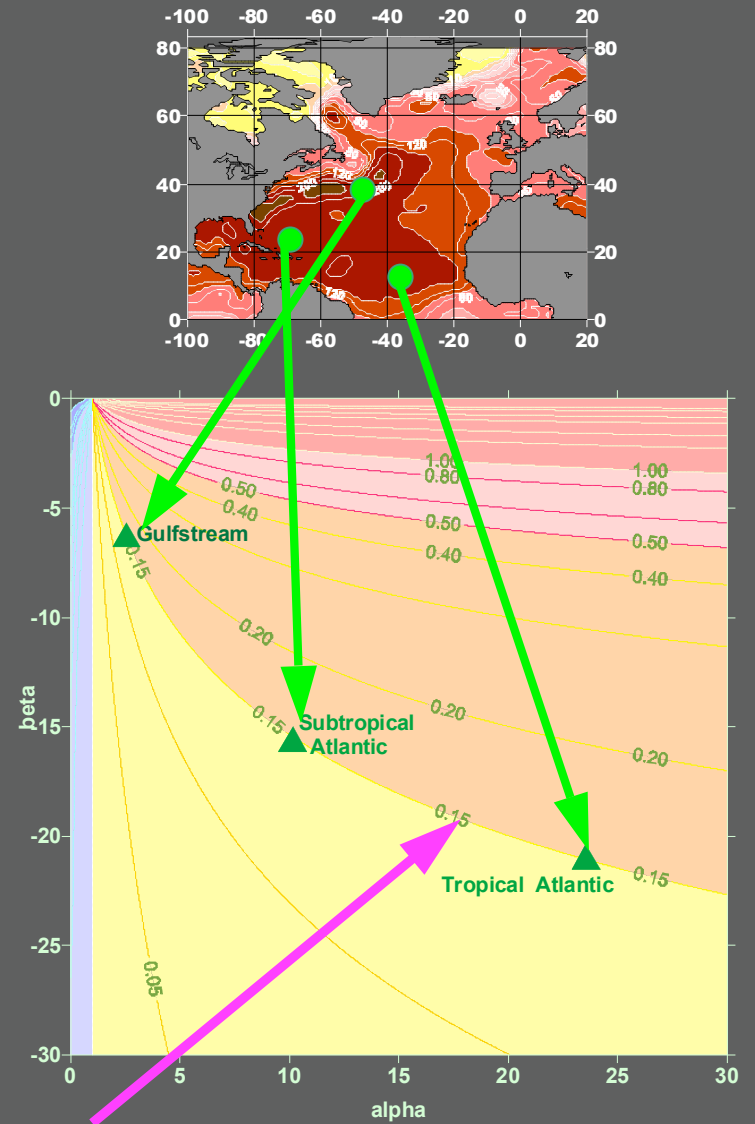
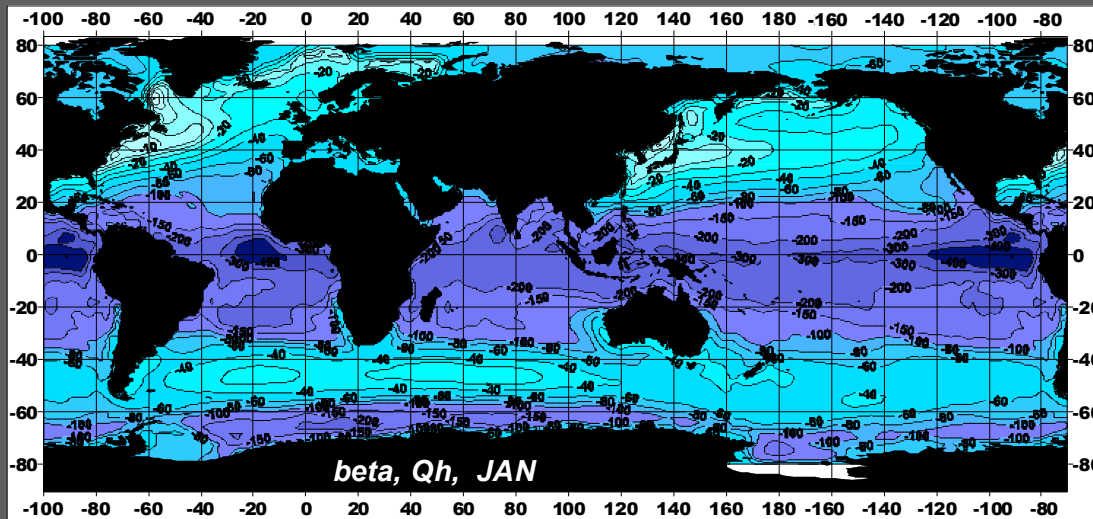
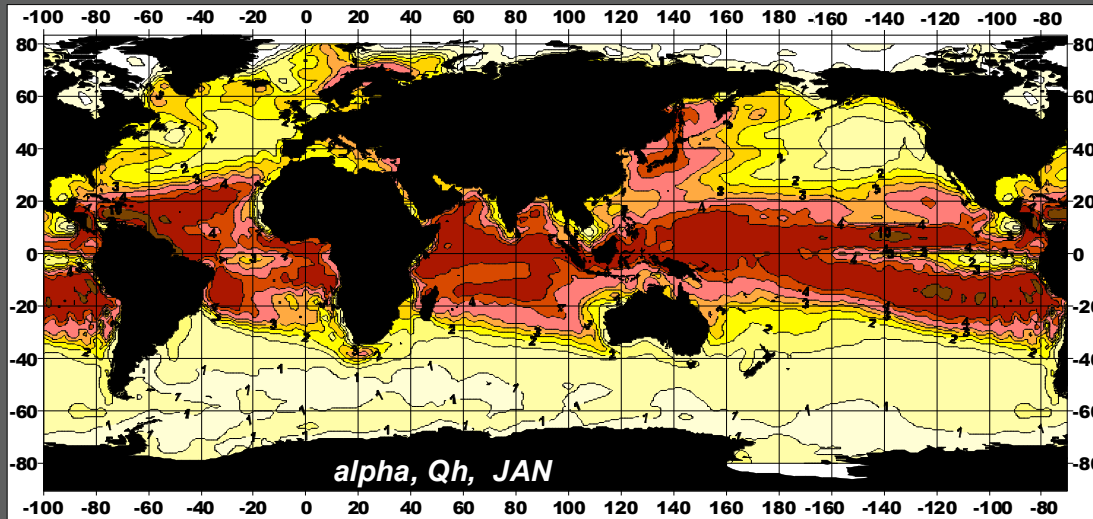
$$\bar{x}_{\text{mod}} = \frac{\ln(1/\alpha)}{\beta}$$



$$\frac{n}{\beta} + \sum_{i=1}^n x_i - n / \sum_{i=1}^n e^{\beta x_i} \cdot \sum_{i=1}^n x_i e^{\beta x_i} = 0; \quad \alpha = n / \sum_{i=1}^n e^{\beta x_i}$$

Parameter estimation:

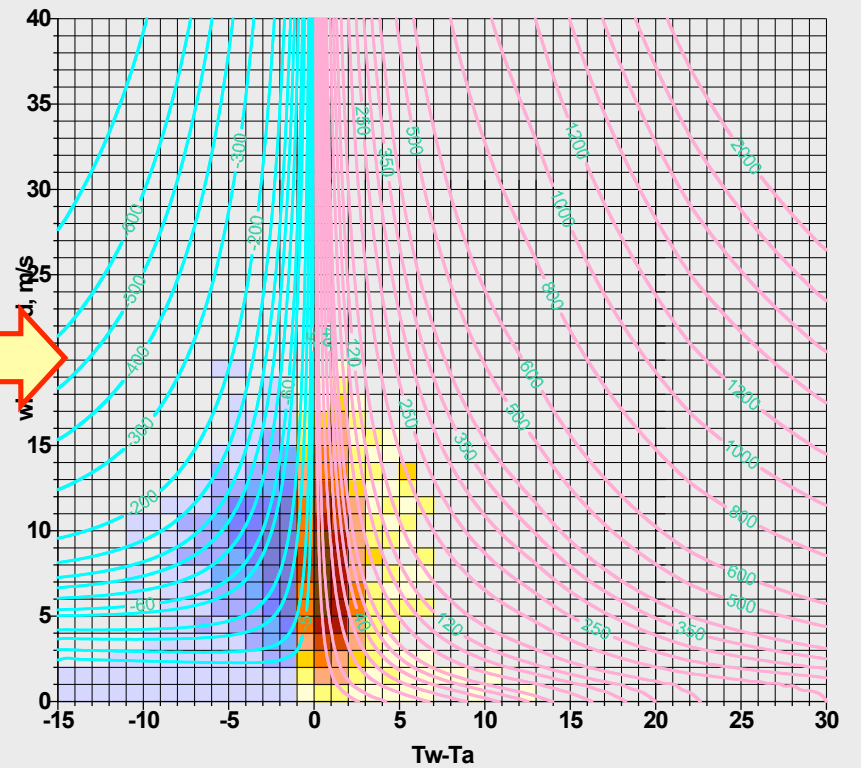
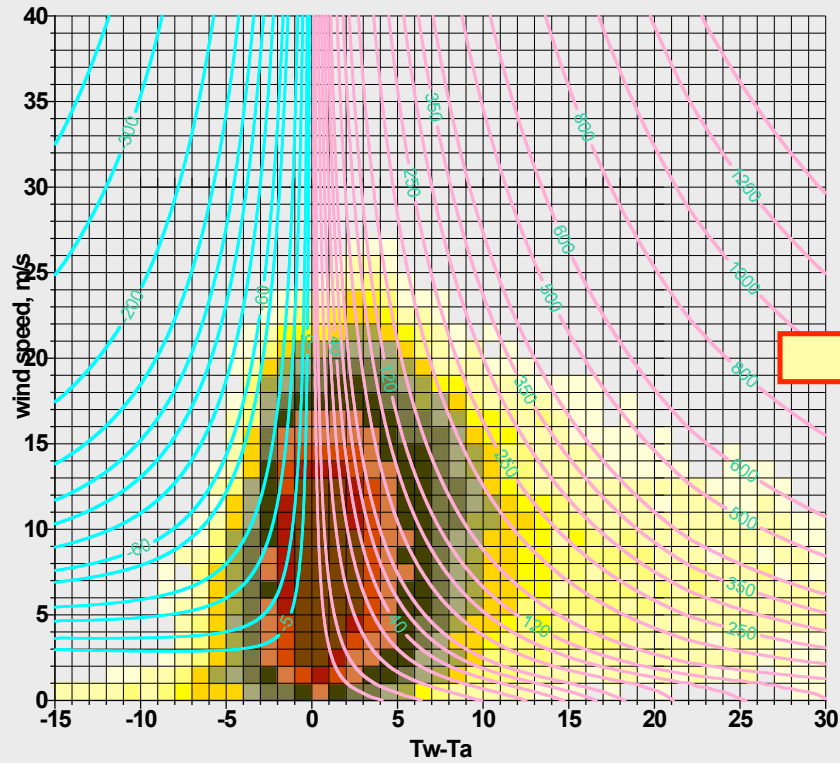
Accounting for different statistical properties of qualitatively similar fluxes



2D-case: Weibull + 2ePDF

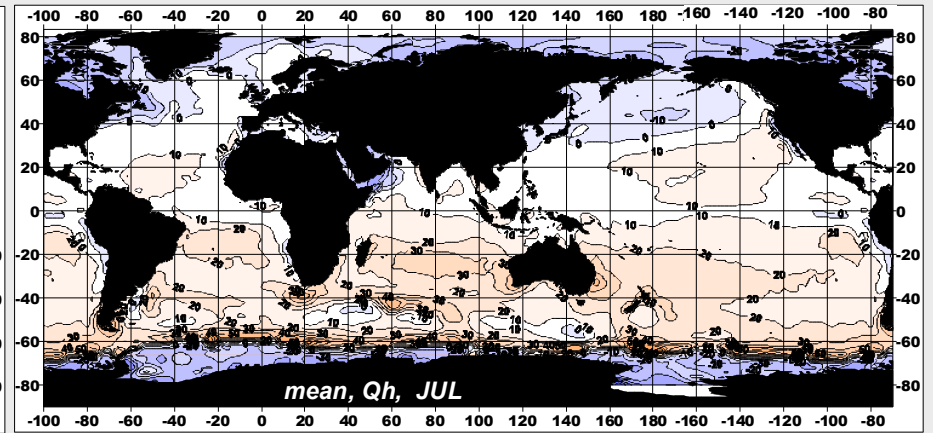
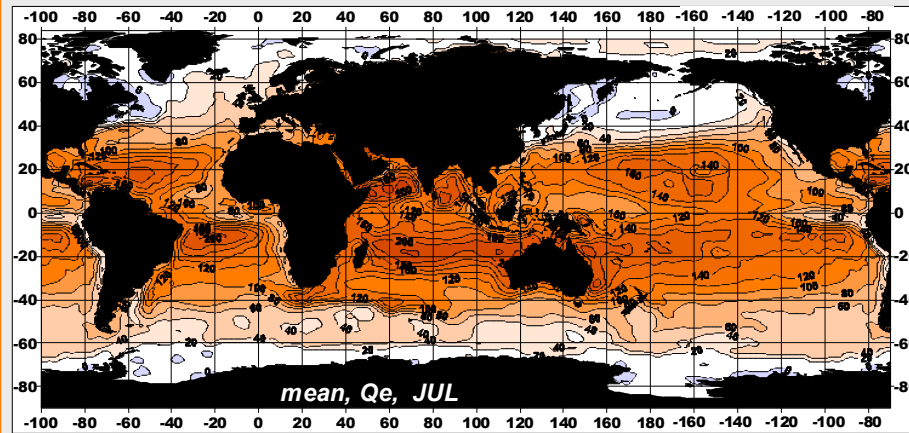
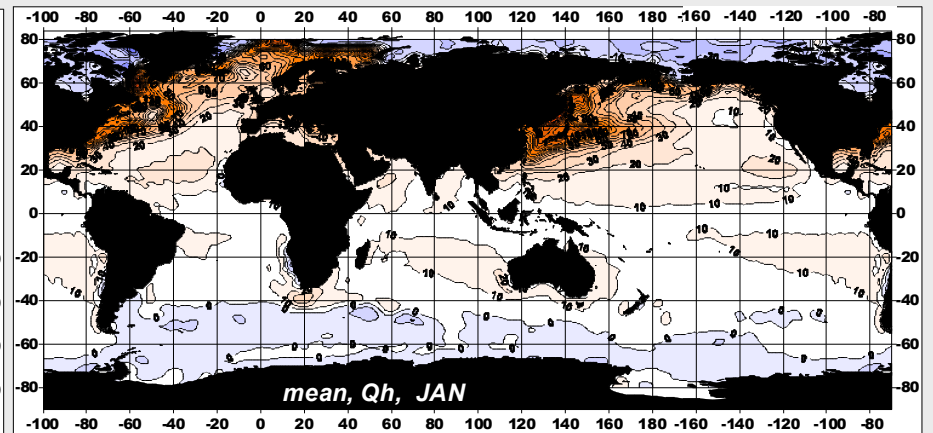
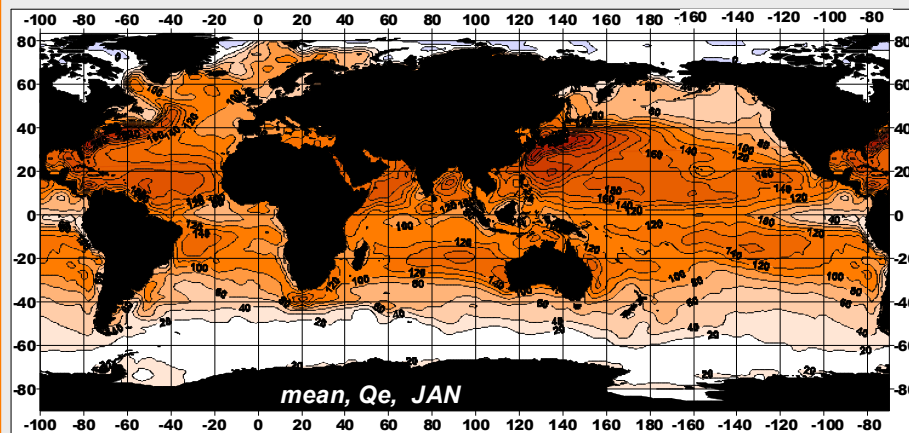
$$P(V | \delta T) \cdot P(\delta T) = \frac{\alpha_V}{\beta_V} \left(\frac{V}{\beta_V} \right)^{\alpha_V - 1} \cdot e^{-\left(\frac{V}{\beta_V} \right)^{\alpha_V}} \cdot (\alpha_T \cdot \beta_T) \cdot e^{\beta_T \delta T} \cdot e^{-\alpha_T \cdot e^{\beta_T \delta T}}$$

WIND - Weibull PDF

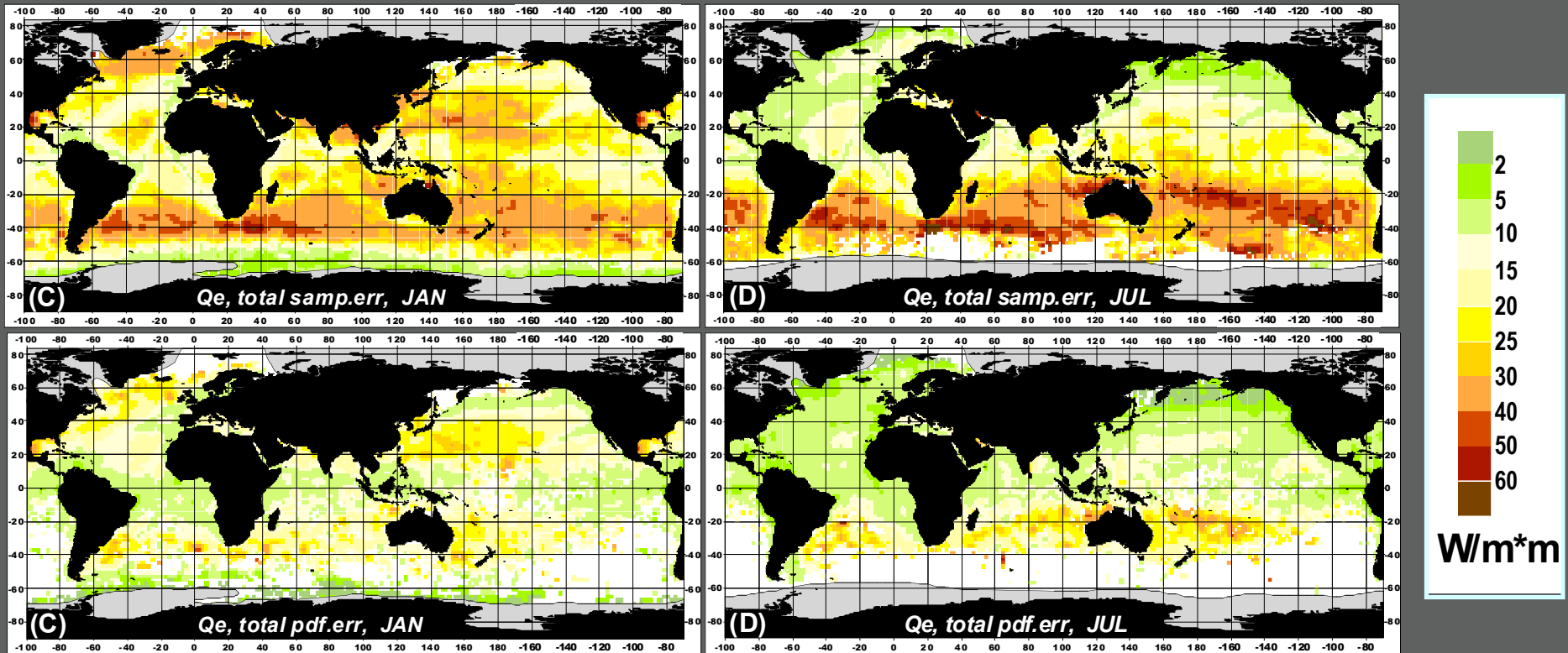


SST-T_{air} -2ePDF

New climatology of turbulent fluxes



Minimization of sampling error



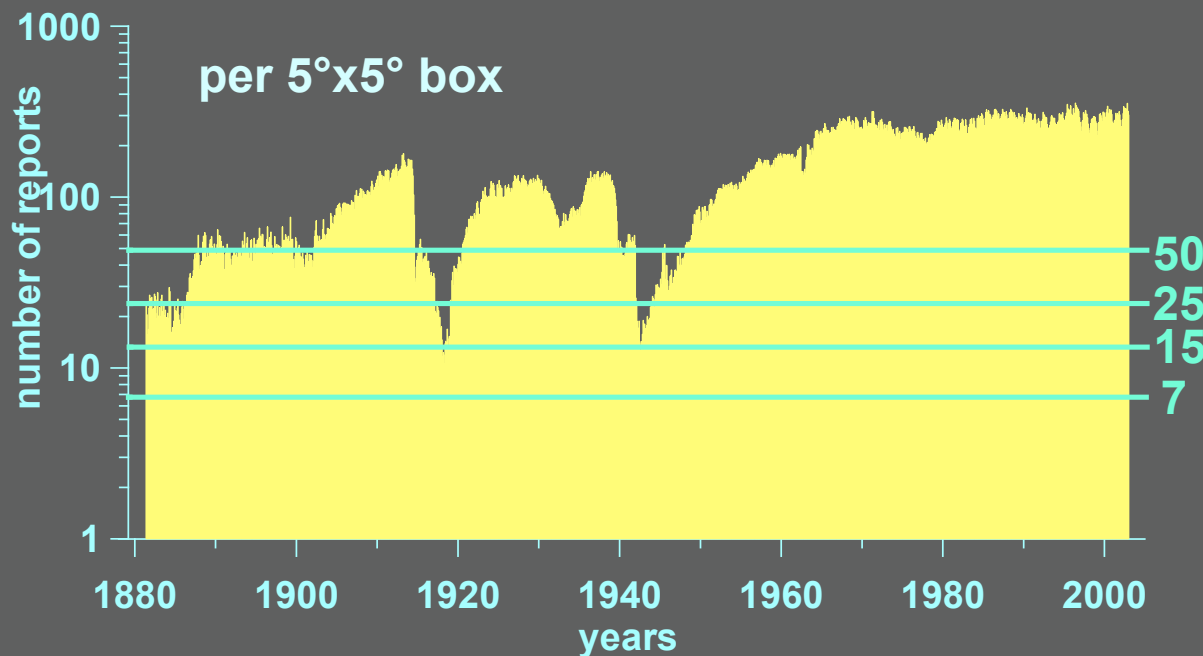
Application of 2eWPDF reduces the sampling uncertainty in 2-10 times (7 times on average)

Task: To derive homogeneous time series of turbulent fluxes in the North Atlantic

What means “homogeneous”?

- Sampling in 1960s+ should be as bad as before WW2
- Impact of parameterizations should be minimized
- Impact of changes in obs. practices should be minimized

The fluxes will not be correct, but variability might be reliable



MC-sub-sampling
for $n=7, 15, 25, 50$
per 5°x5° box

Technology of the reconstruction

Parameterizations:

- COARE-3.0 (with some simplifications)

Parameters:

- Wind – Beaufort only, WMO1100 => to Lindau (1995) scale (not after 1986)
- SST – buckets only, no engine intakes (not after 1986)
- Air temperature – all
- Relative humidity – multi-regression reconstruction

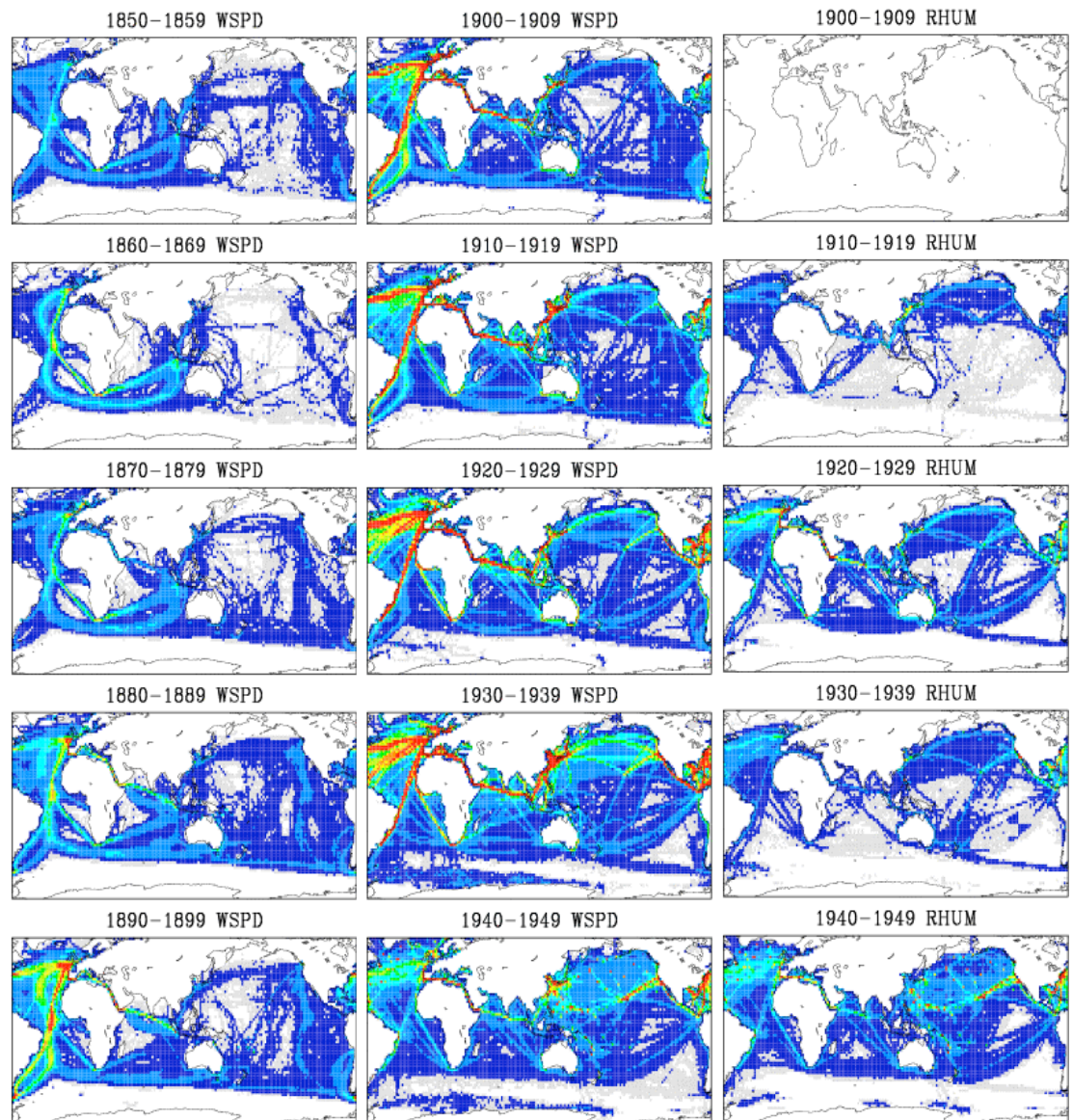
Production – iterative run for every month:

- 1st guess - 2eWPDF derivation of monthly statistics for 10° boxes south of 40°N and for gerrymander cells north of 40°N
- 2eWPDF computation for $5^\circ \times 5^\circ$ boxes, if the PDF does not fit at 95% level, use the 1st guess
- then – repeat the procedure again

Reconstruction of humidity

Sampling of humidity:

- no humidity before 1910
- humidity is sparse before 1950

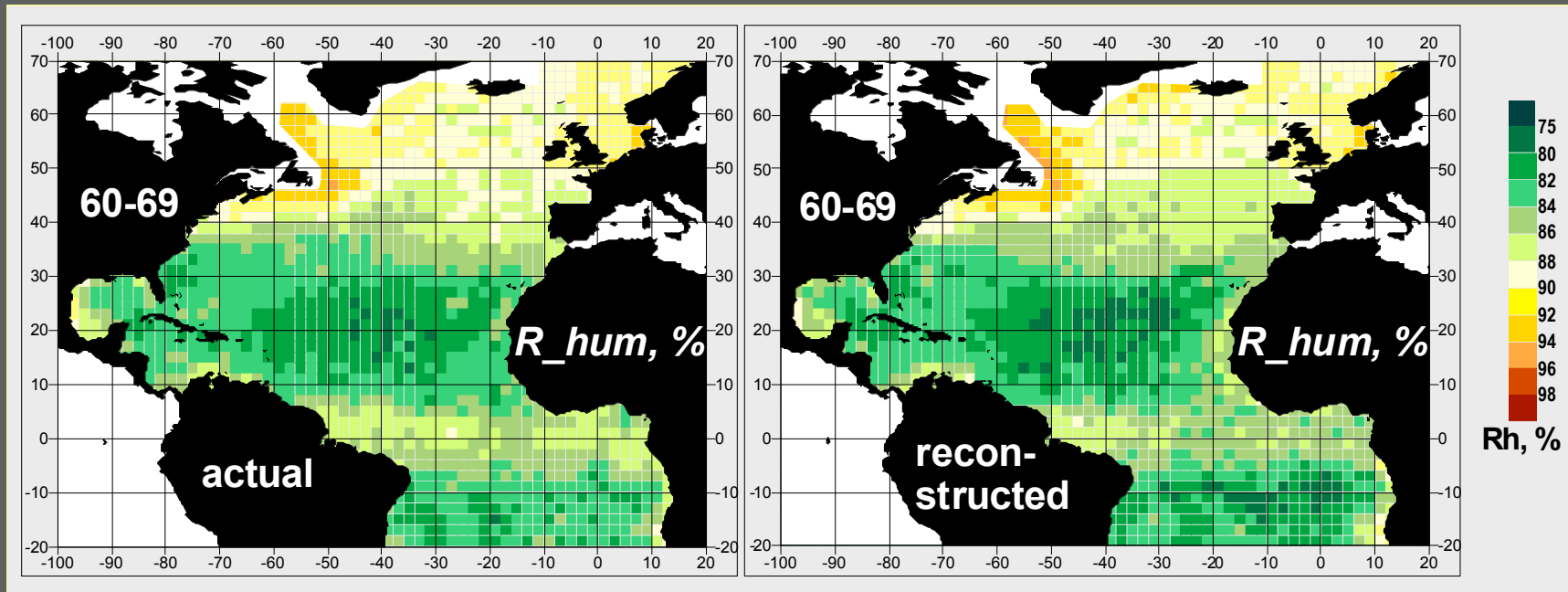


Woodruff et al. 2005

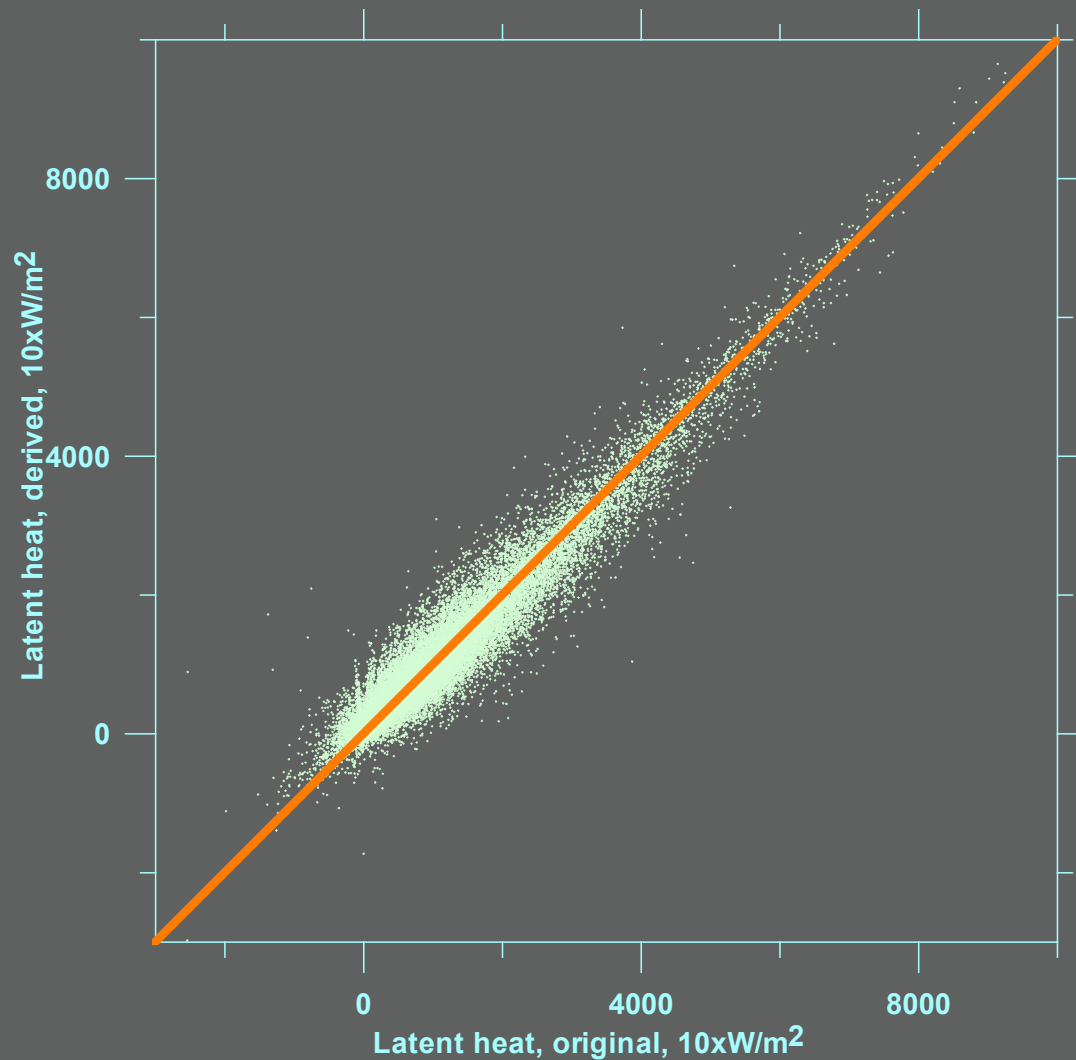
Reconstruction of relative humidity

Strategy:

- Statistical multiregression approach (not a neural network yet) for the decade of 1960s
- Deviations from seasonal climatology are considered
- Exponential functions + polynomials for T_a , V , SLP
- Done locally, for every month and every box



Effect of humidity in latent heat flux

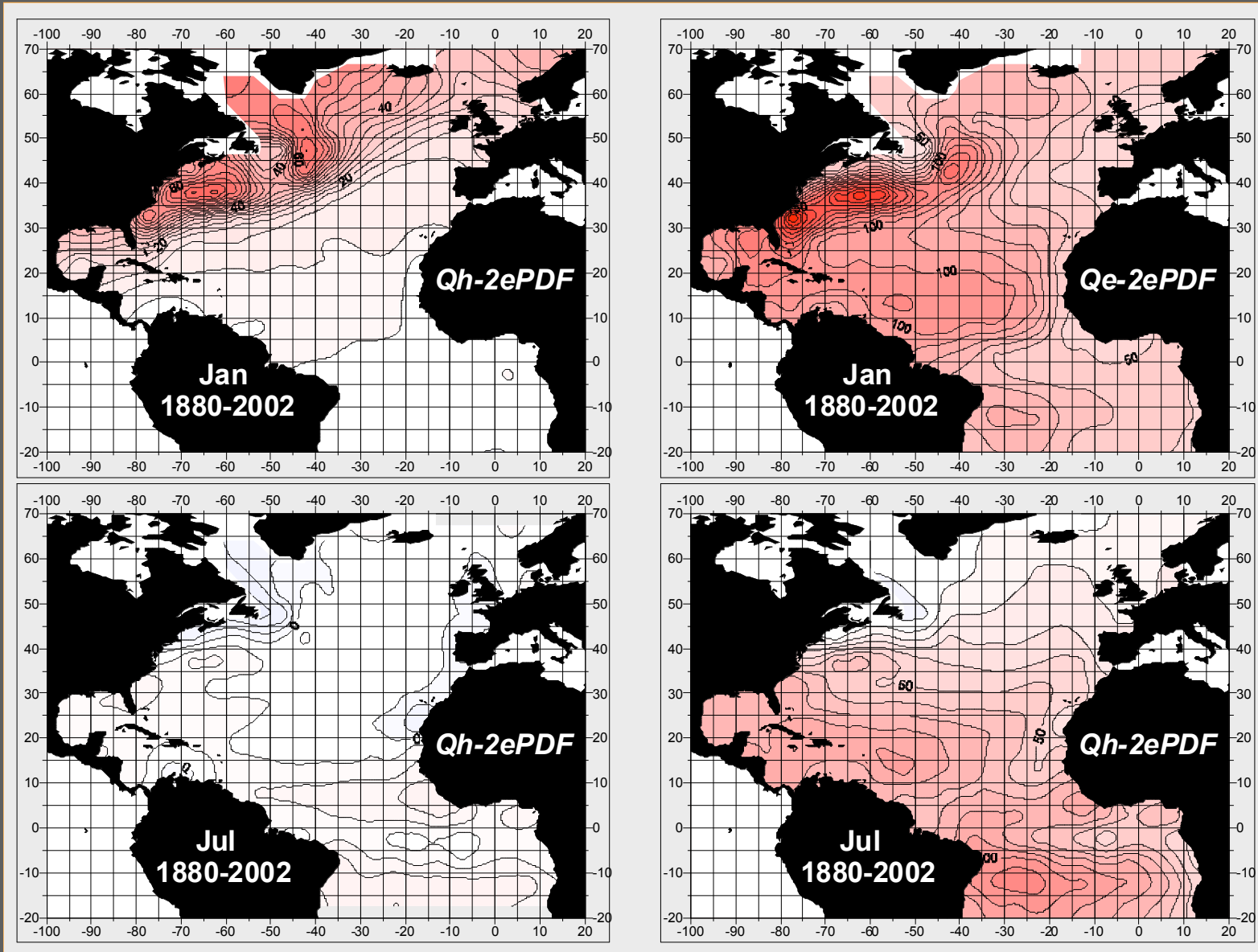


RMS = 7 W/m²

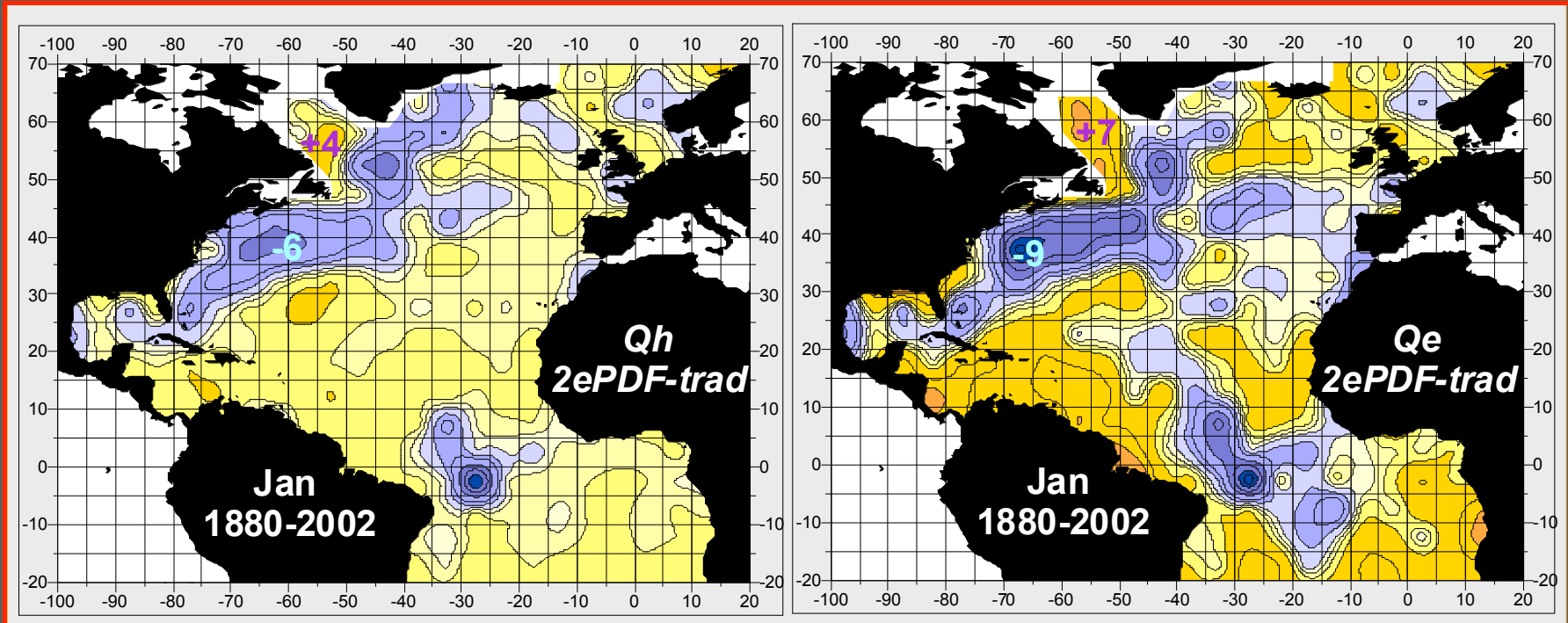
Slope = 0.996

**Intercept =
+1.4 W/m²**

125-yr (1980-2004) climatology



How it is different from the raw averaging?

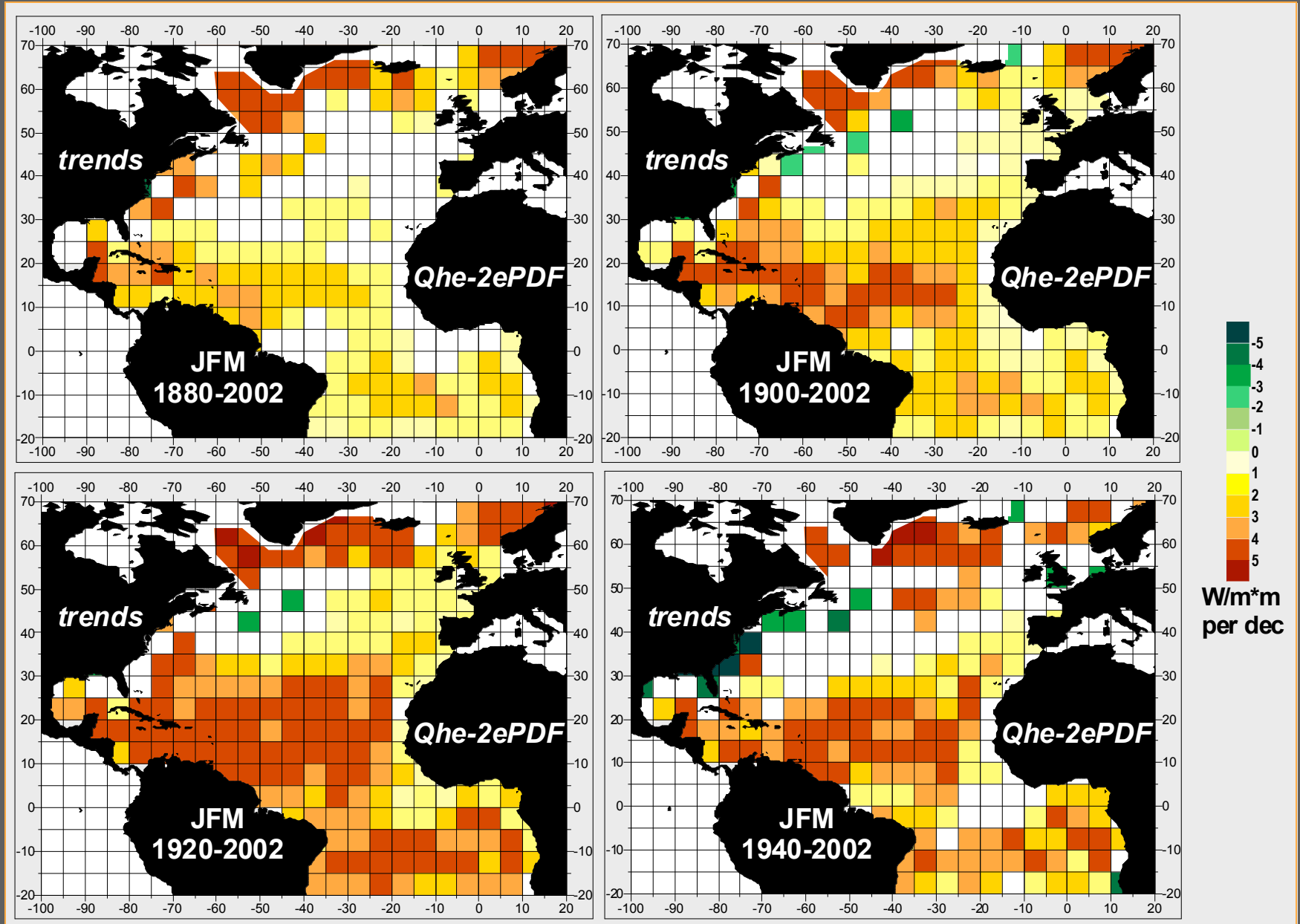


Differences

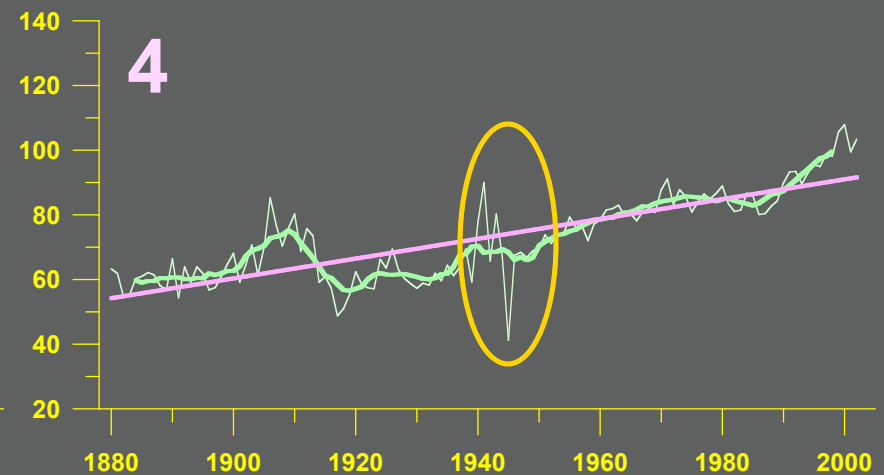
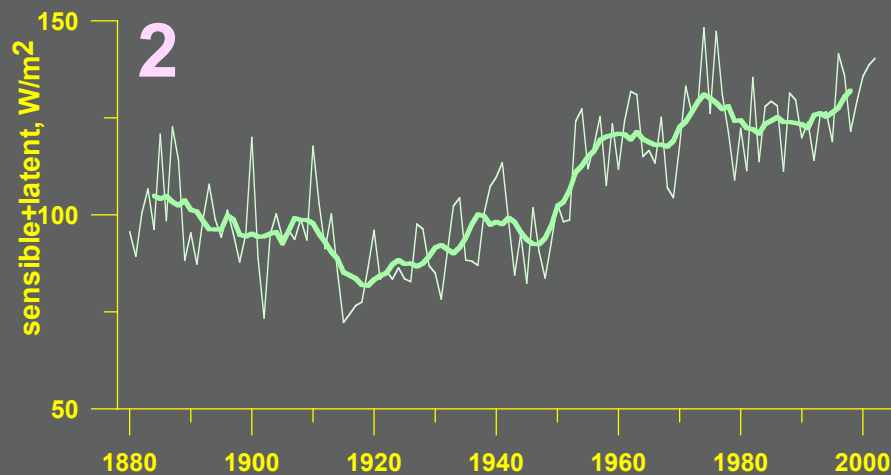
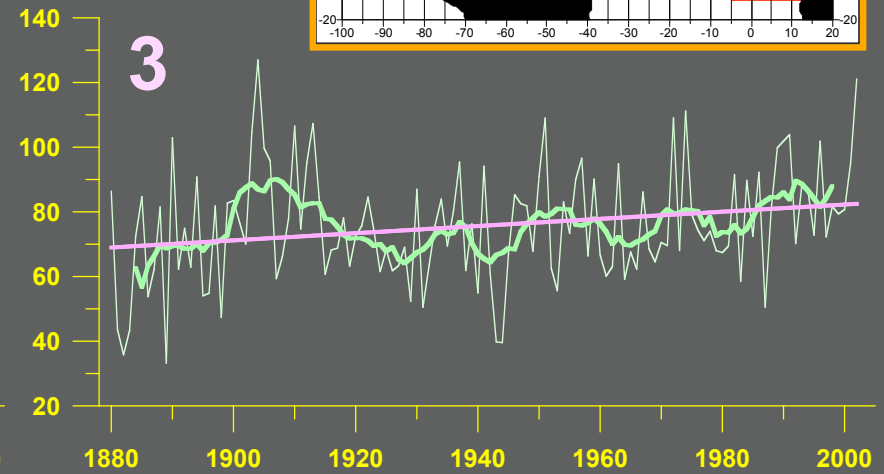
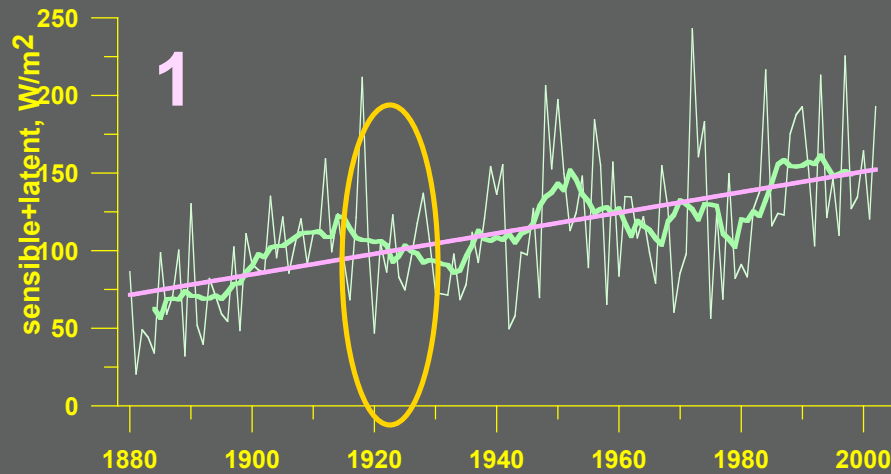
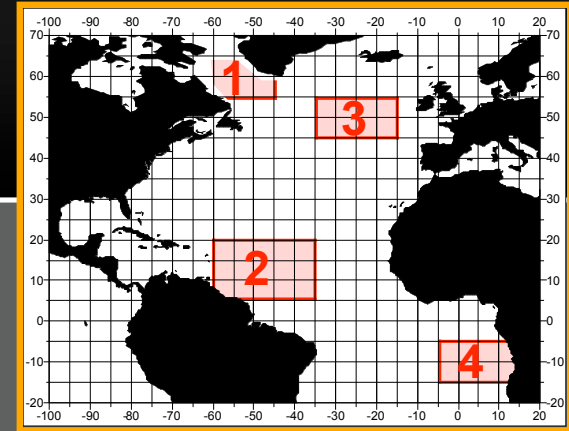
“2eWPDF-derived minus traditionally averaged”
fluxes:

Sampling effect is clearly visible

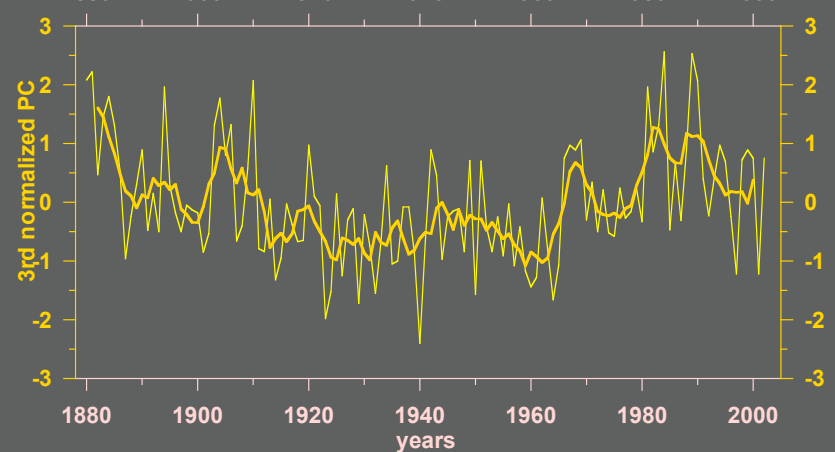
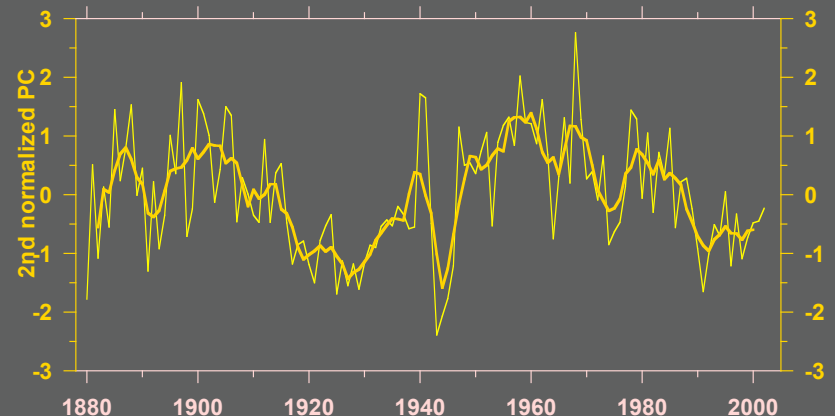
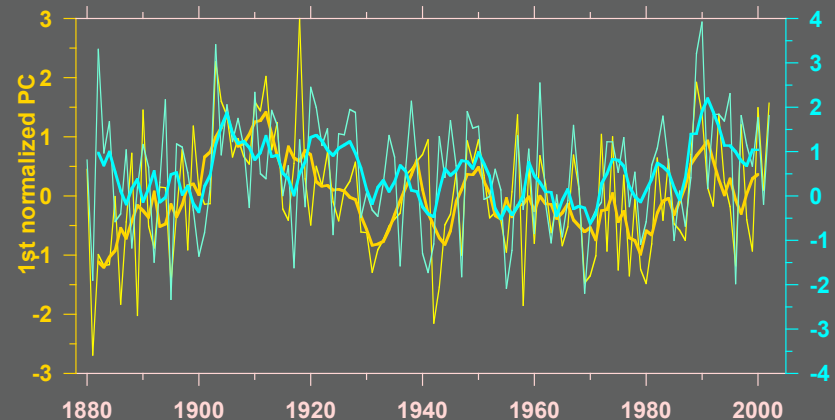
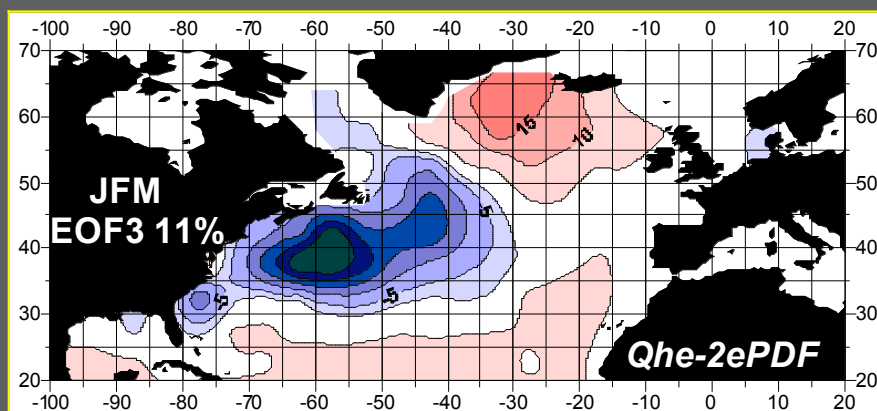
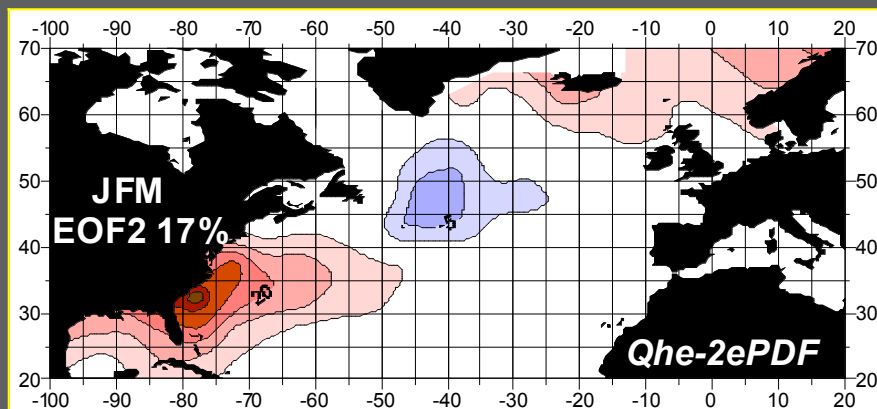
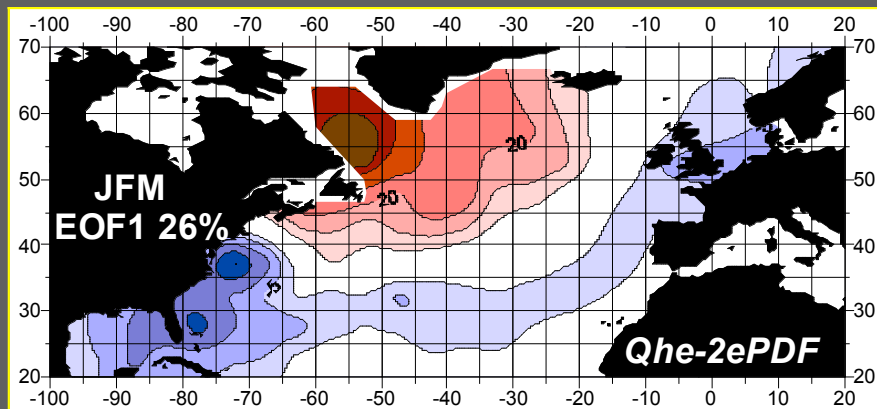
Winter linear trends



Regional time series of winter sensible+latent heat flux

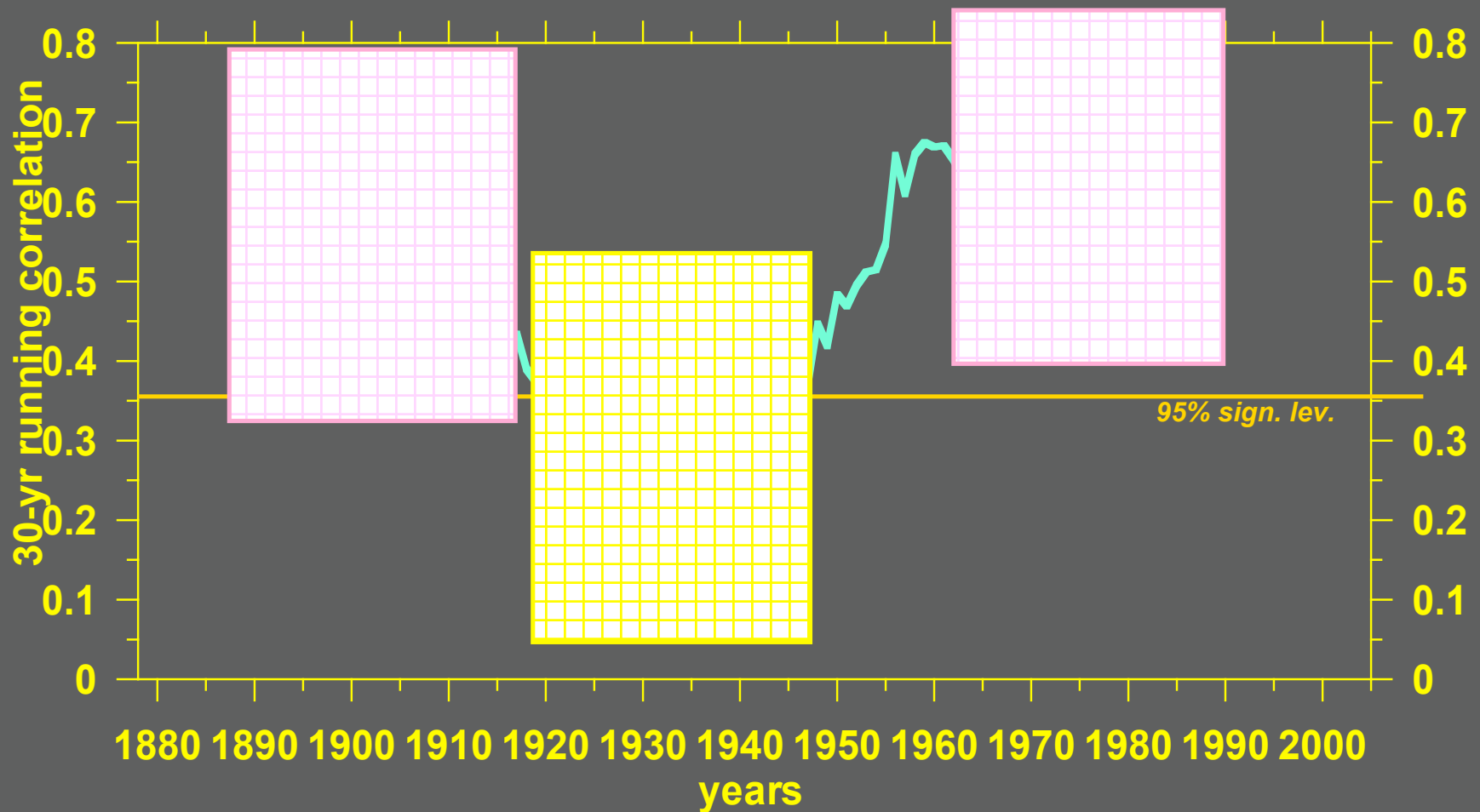


Winter (JFM) EOFs of sensible+latent flux

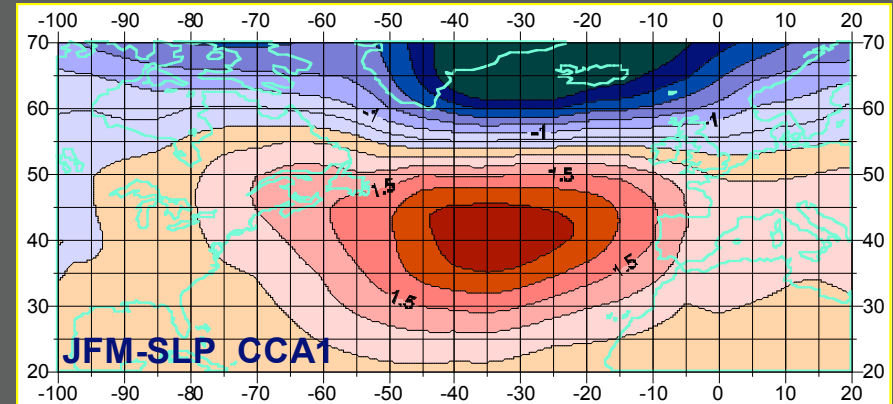
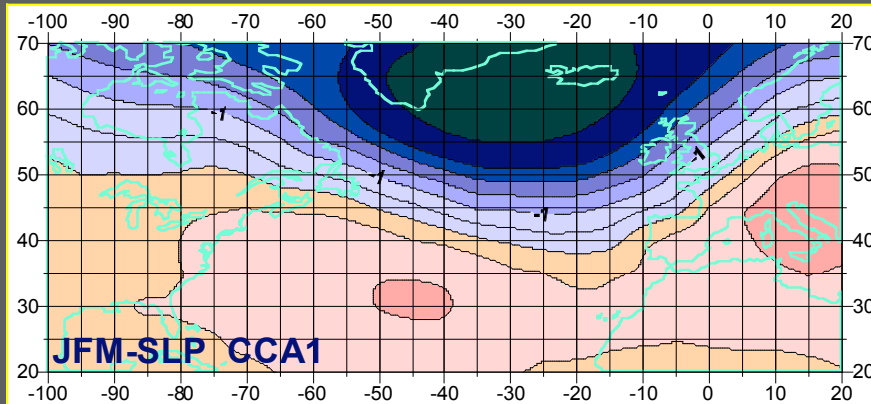
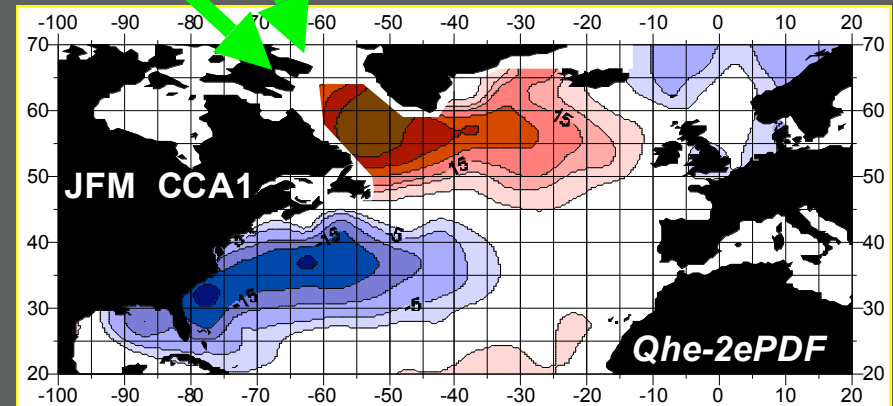
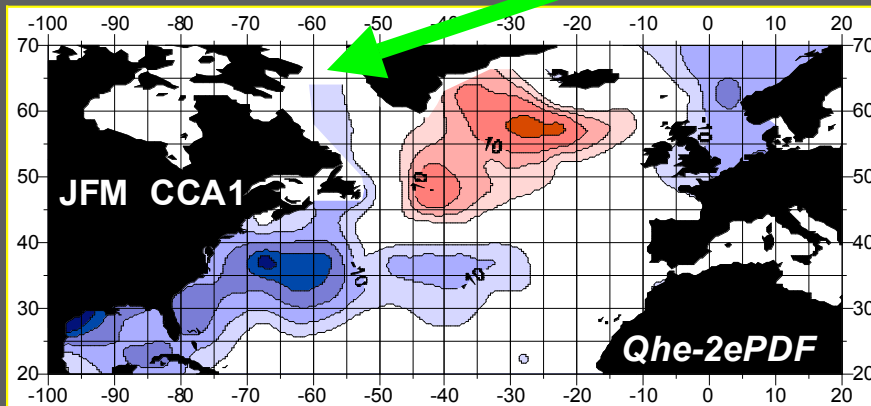
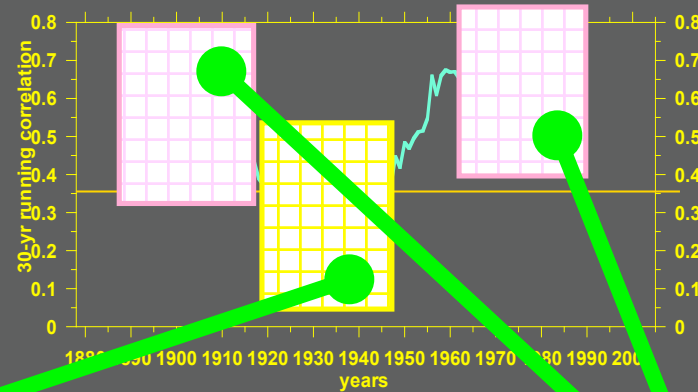


Link of the leading mode of turbulent fluxes to NAO

30-yr running correlation changes over time considerably



Canonical correlation of the heat fluxes with HadSLP for different periods



Conclusions

- ★ **2eW-PDF allows for the minimization of sampling biases in turbulent fluxes from 2 to 7 times and Allows for the development of long-term homogeneous flux time series**
- ★ **5-degree 125-yr homogeneous time series of turbulent fluxes (1880-2004) were derived – multi-decadal variability is visible**
- ★ **During 1915/20 – 1950/55 winter surface turbulent fluxes are loosely connected with NAO, being closely related to NAO during the decades before and after this period**

Special thanks

Steve Worley, DSS, NCAR, Boulder

Scott Woodruff, CDC, NOAA, Boulder

**for many years of reliable feedback in
ICOADS data management**

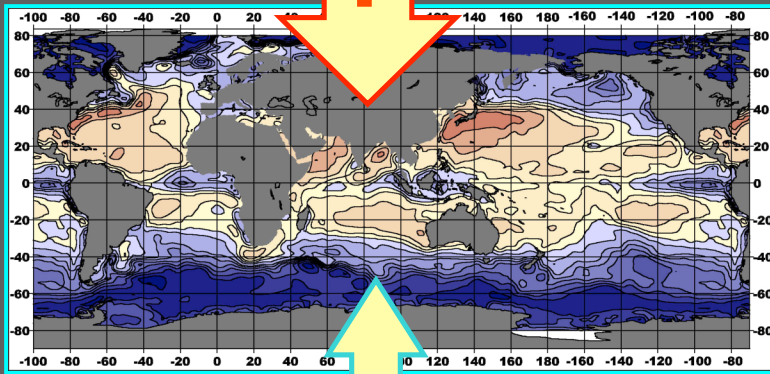


VOS-based ocean-atmosphere fluxes



Measurements and variable corrections:

SST, T_a , q , SLP, V , cloud cover,
wave heights and periods

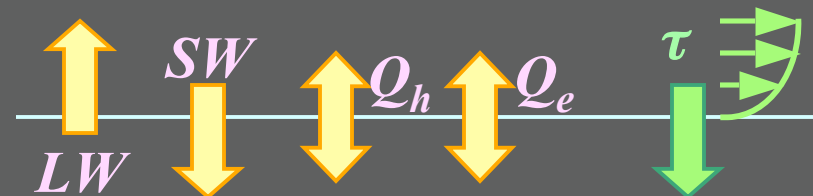


Objective analysis
and averaging:

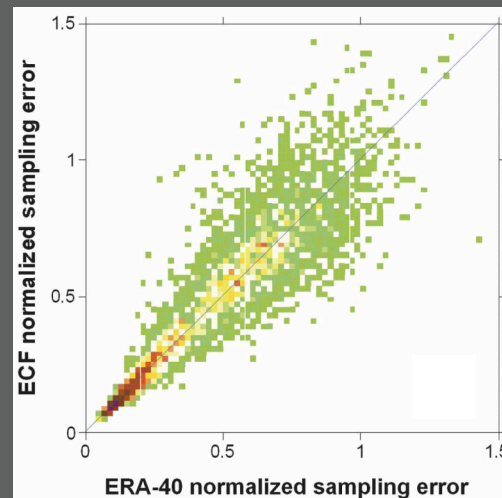
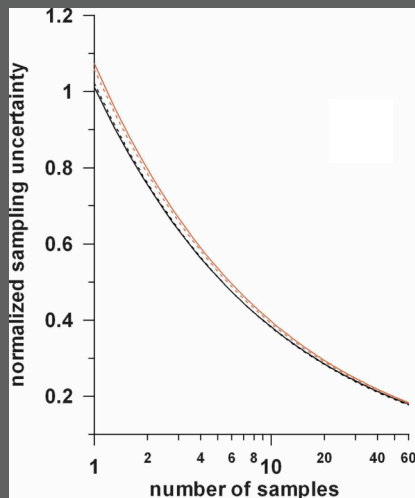
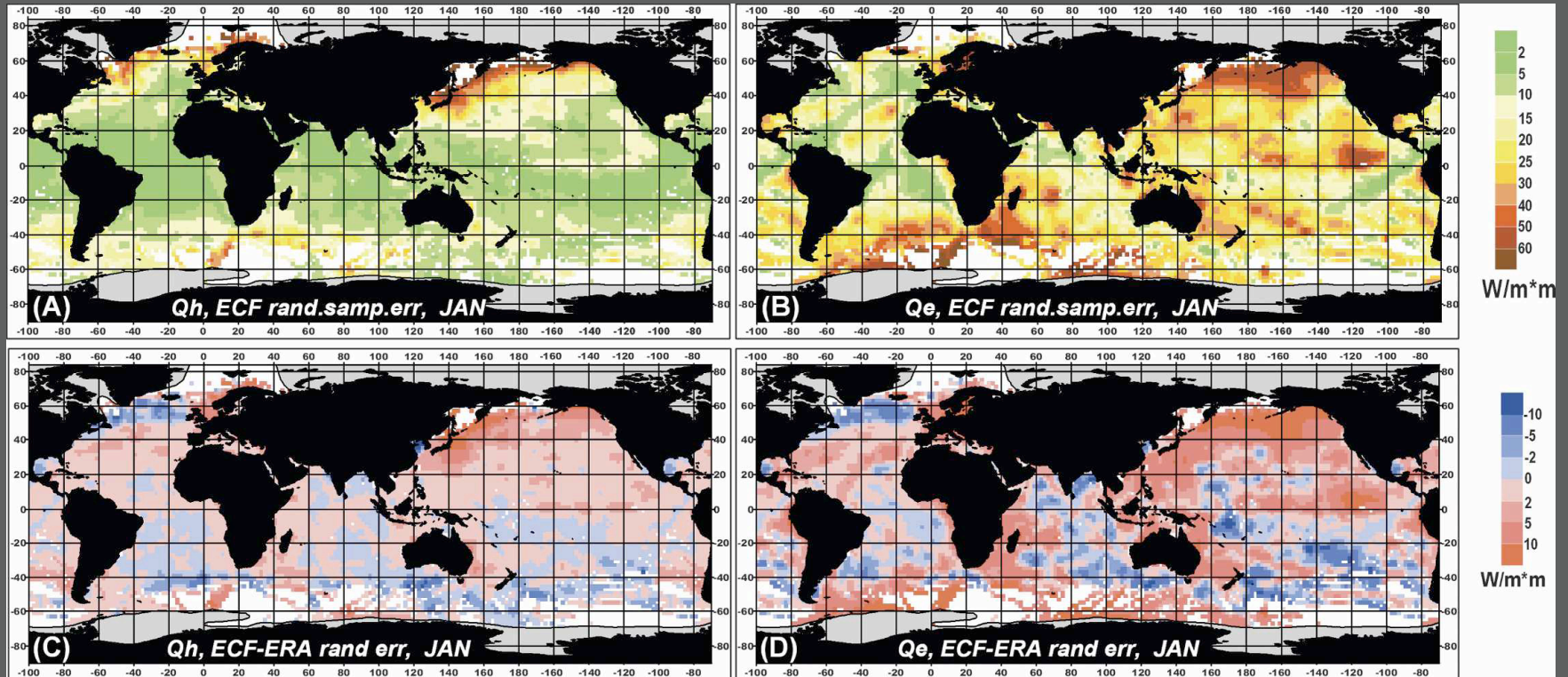
Global air-sea
flux fields

Parameterizations:

$$\begin{aligned}
 LW &= \varepsilon \sigma SST^4 F_{lw}(C_n, T_a, e_z), \\
 SW &= Q_o F_{sw}(C_n, T_a, e_z, h_o)(1-\alpha), \\
 Q_h &= C_p C_t \rho (SST - T_a \gamma) V, \\
 Q_e &= L C_e \rho (0.622/P) (e_o - e_z) V, \\
 H &= SW - LW - Q_h - Q_e \\
 \tau &= \rho C_v (V, h_w) V^2
 \end{aligned}$$

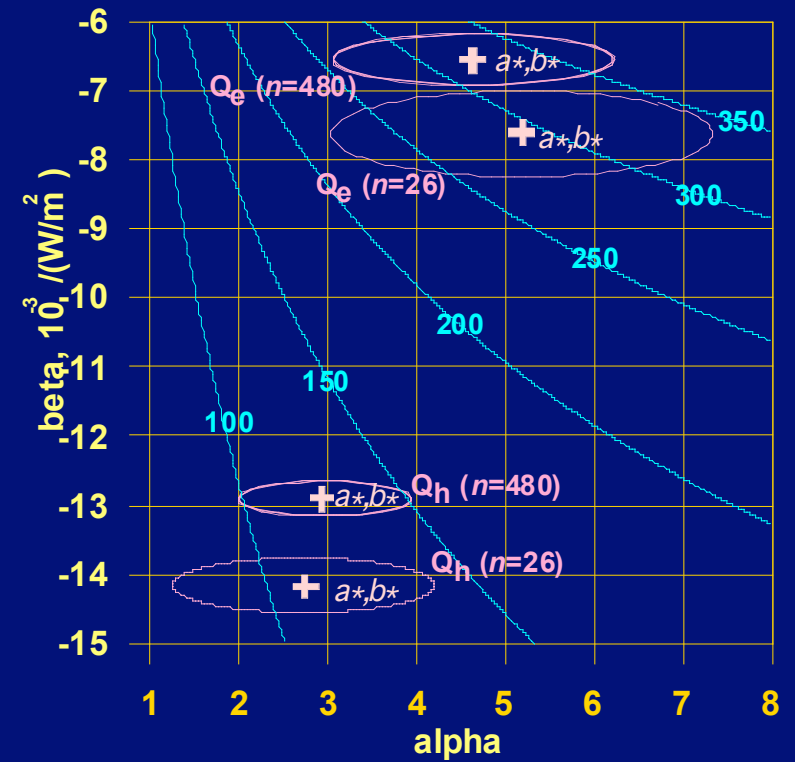
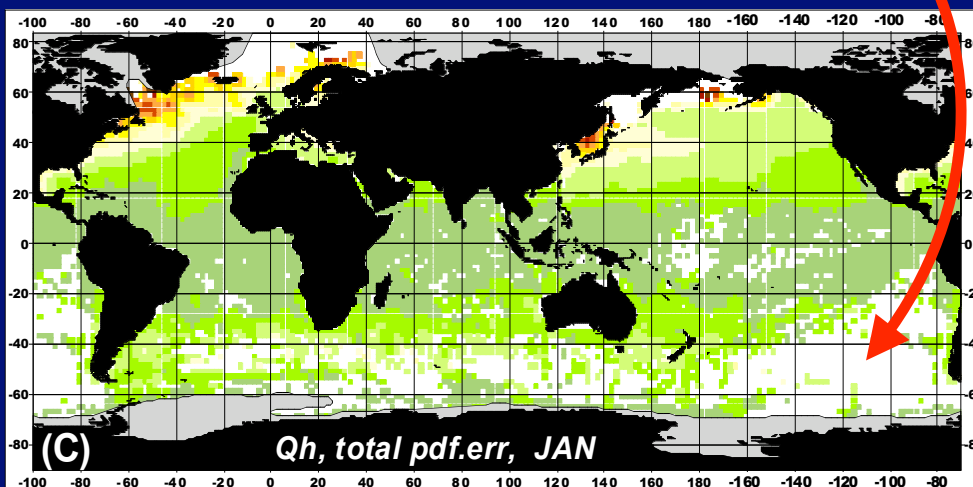
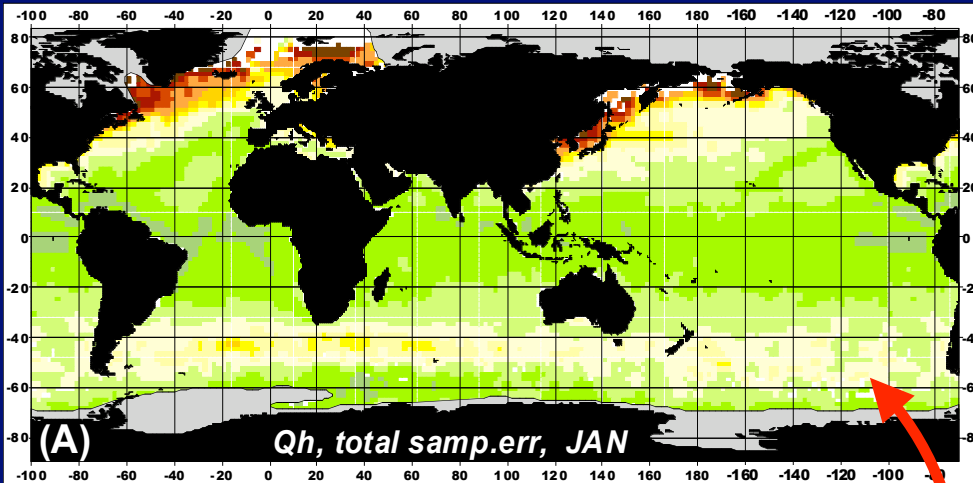


Sampling errors from different NWP

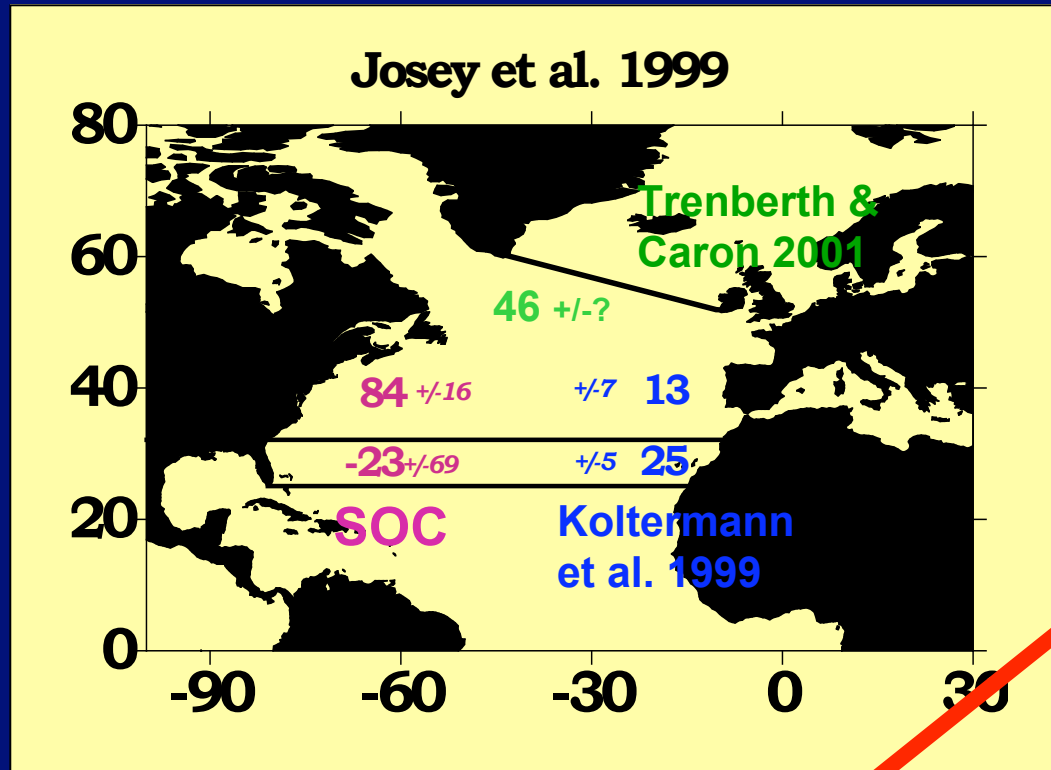


Actual values may be different, but when scaled with *std* sampling error is practically equal, when derived from different NWP

$$\frac{(\alpha - \alpha_*)^2}{\sigma_1^2} - \frac{2\rho}{\sigma_1\sigma_2}(\alpha - \alpha_*)(\beta - \beta_*) + \frac{(\beta - \beta_*)^2}{\sigma_2^2} = 6(1 - \rho^2)$$

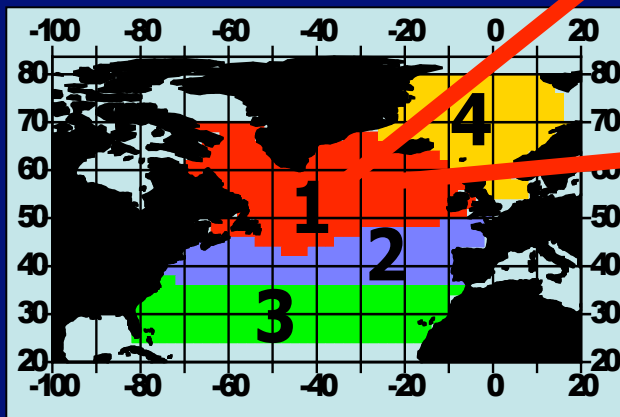


Estimating regionally integrated heat flux using 2e PDF



Uncertainty in the integrated sensible+latent flux

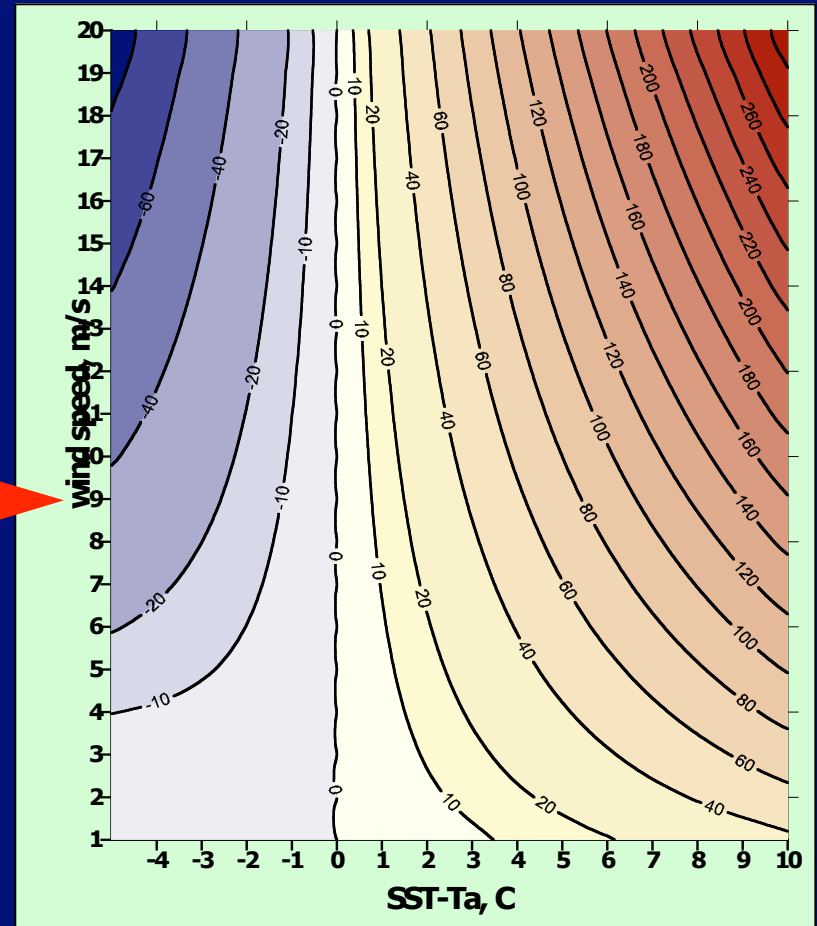
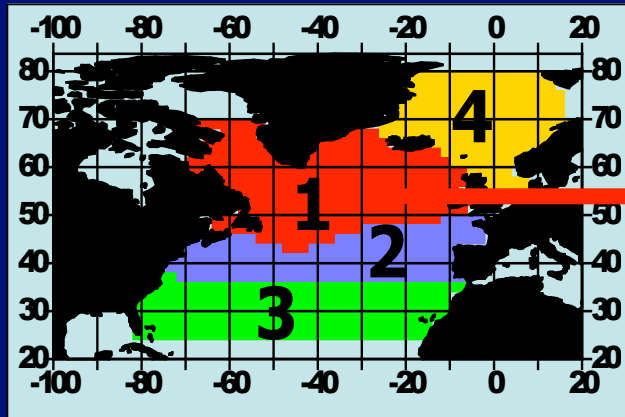
VOS-like sampling:
 $\Delta = 0.67 * 10^{14} \text{ W}$



2e-PDF - reconstruction:
 $\Delta = 0.23 * 10^{14} \text{ W}$

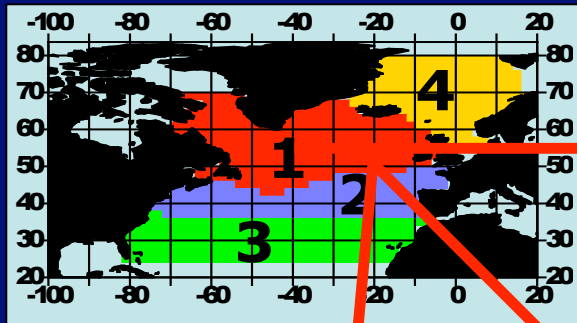
Regional balances: a new domain for integration

$$Q_h = C_p \cdot \rho \cdot C_t(\delta T, V) \cdot \delta T \cdot V$$

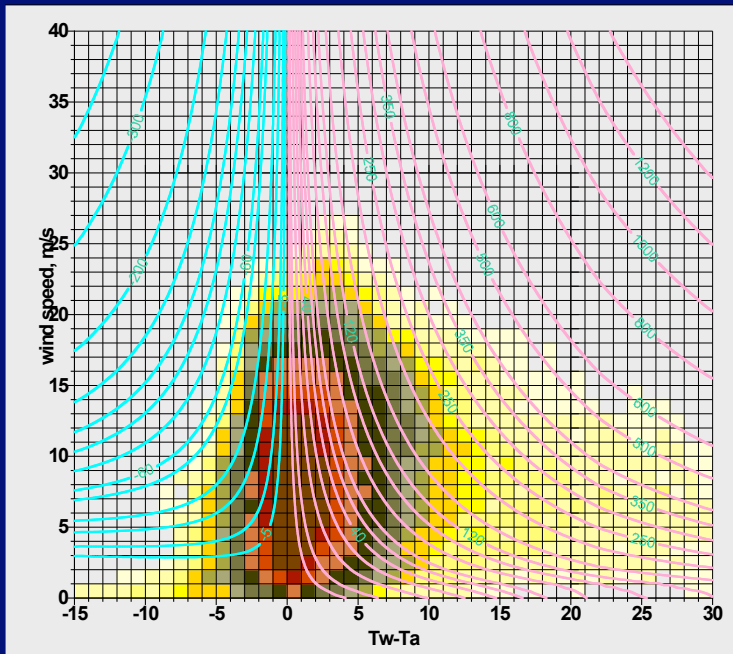


$$Q_{\Sigma} = \int_t dt \int_S Q dS = \int_V dV \int_{\delta T} Q d(\delta T)$$

2D-distributions of turbulent fluxes

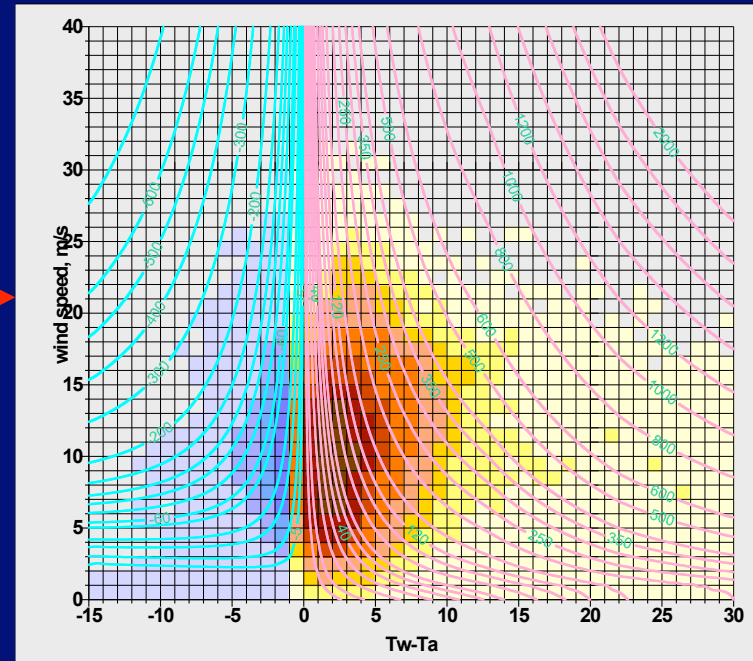


2D- PDF of sensible flux



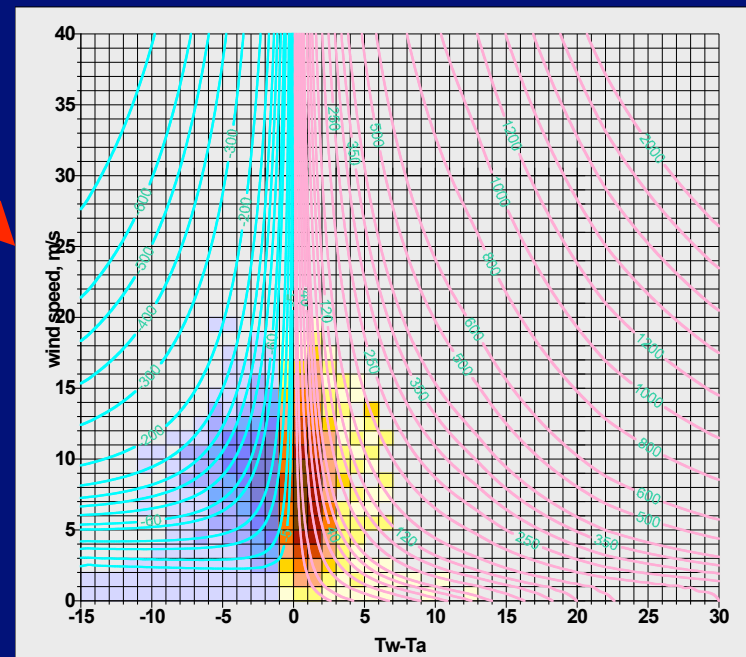
Fluxes accumulated over $\delta T, V$ -classes

Q_e

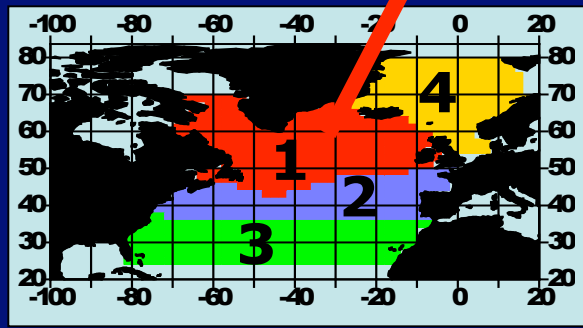
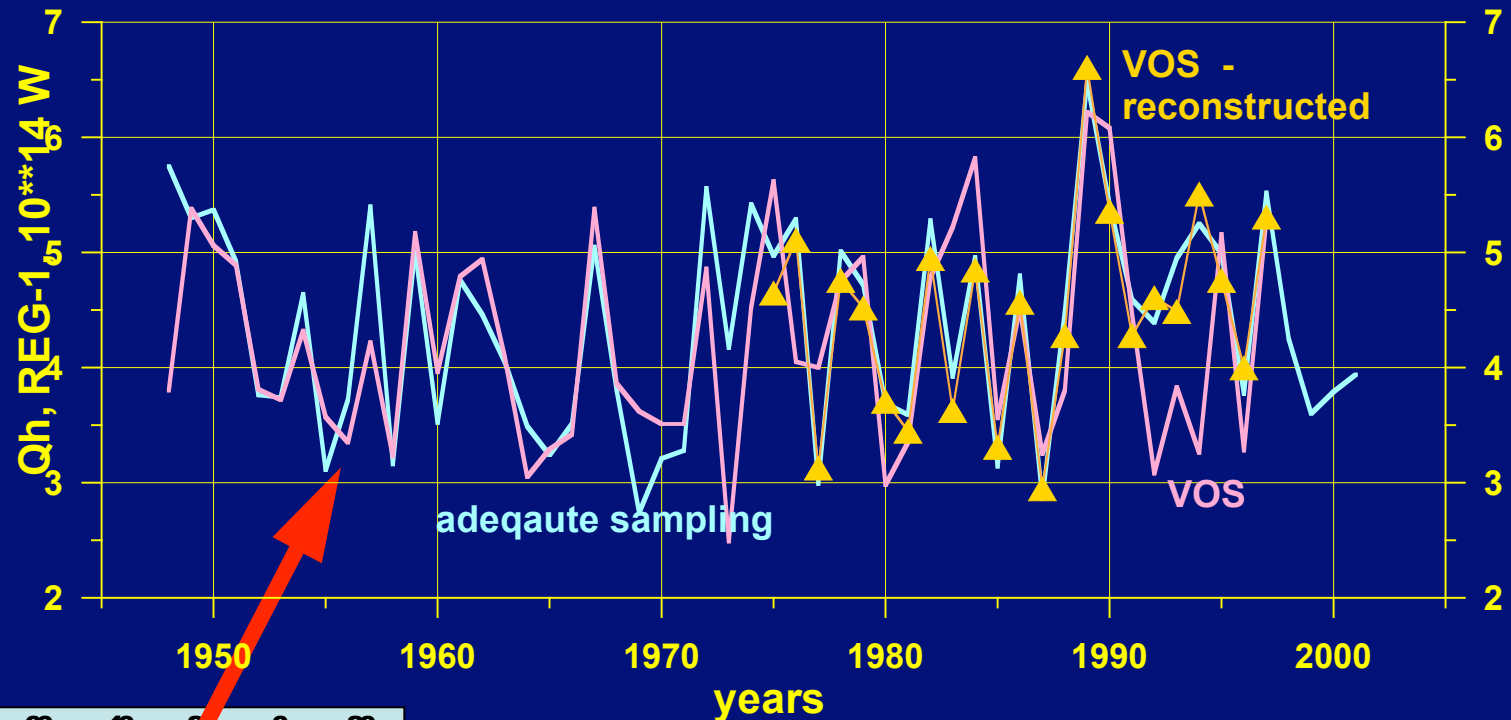


$10^{11} W$

Q_h



Estimating regionally integrated heat flux using two dimensional 2e-Weibull PDF



2e-PDF - reconstruction:
 $\Delta = 0.23 * 10^{14} \text{ W}$

VOS-like sampling:
 $\Delta = 0.67 * 10^{14} \text{ W}$

2D-2eW-PDF - reconstruction:
 $\Delta = 0.11 * 10^{14} \text{ W}$

