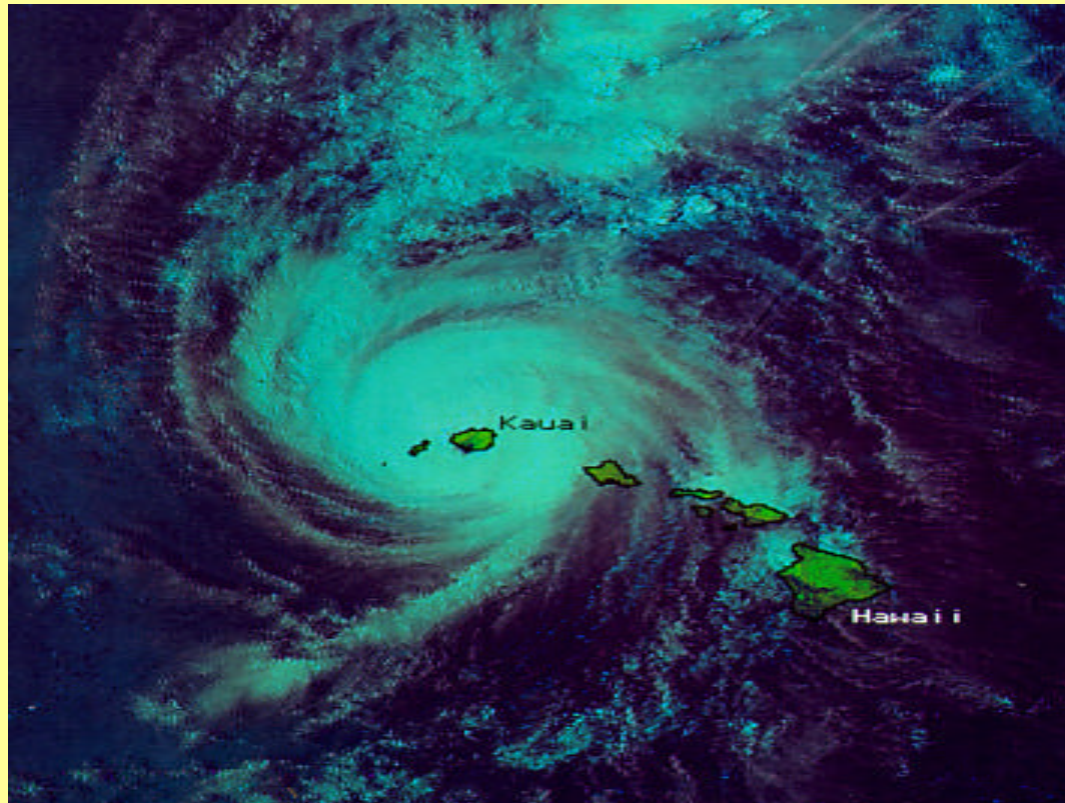


# Hurricane induced Wave-Surges at landfall



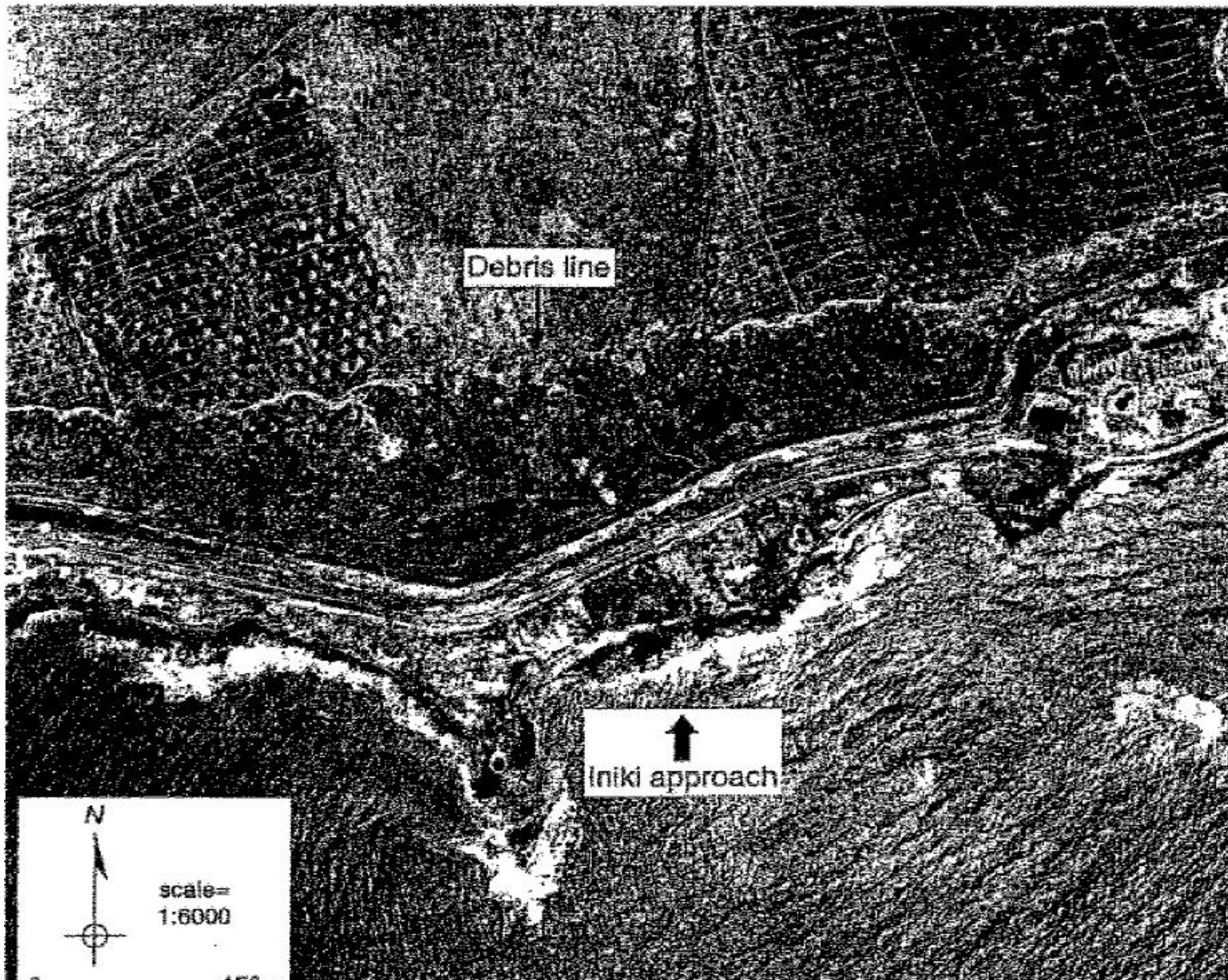
Chung-Sheng Wu

Arthur Taylor, Jye Chen and Wilson Shaffer

Meteorological Development Laboratory

Office of Science and Technology, National Weather Service

National Oceanic and Atmospheric Administration

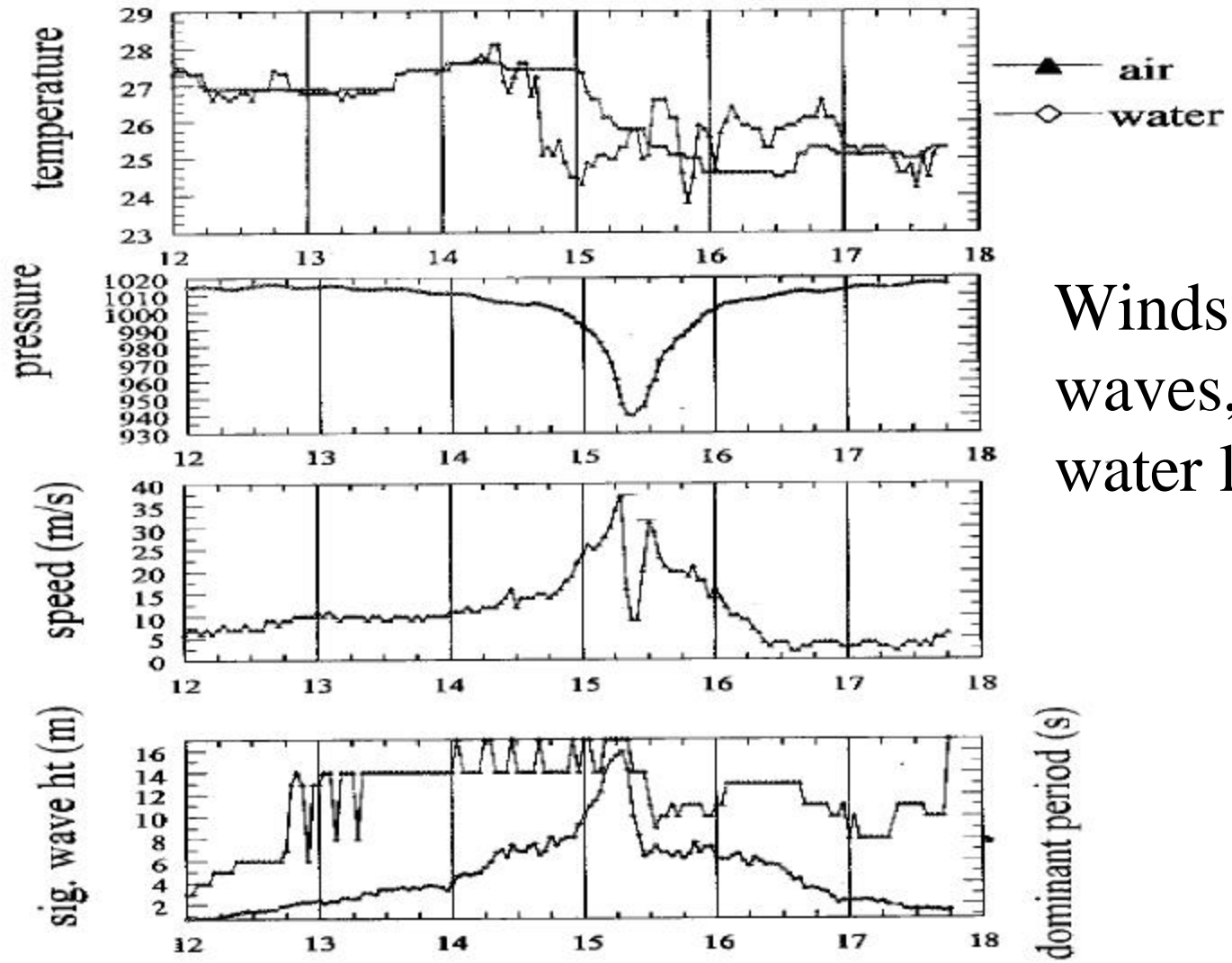


Debris line

↑  
Iniki approach

N  
↑  
○  
scale=  
1:6000

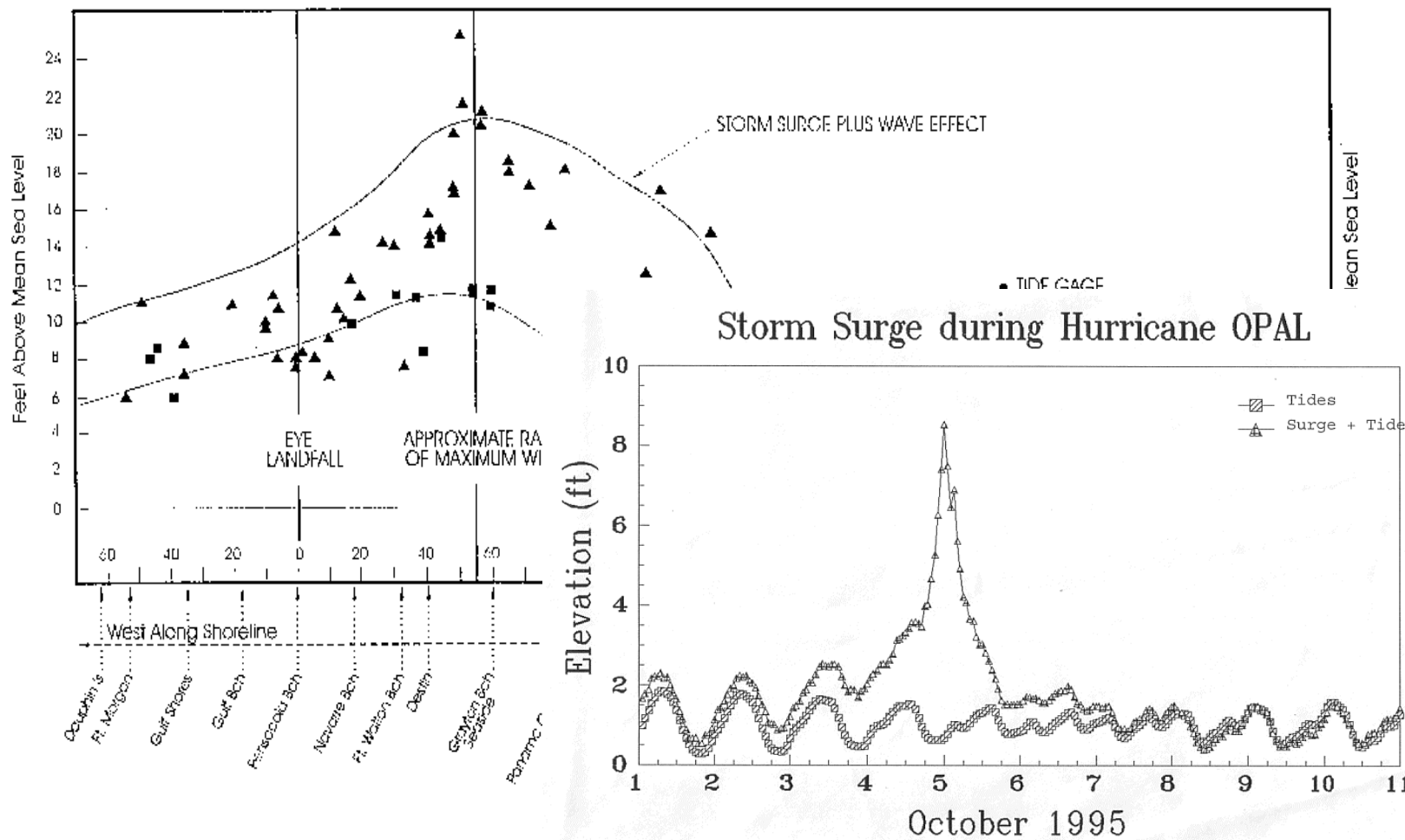
# Buoy 41010, Hurricane Floyd



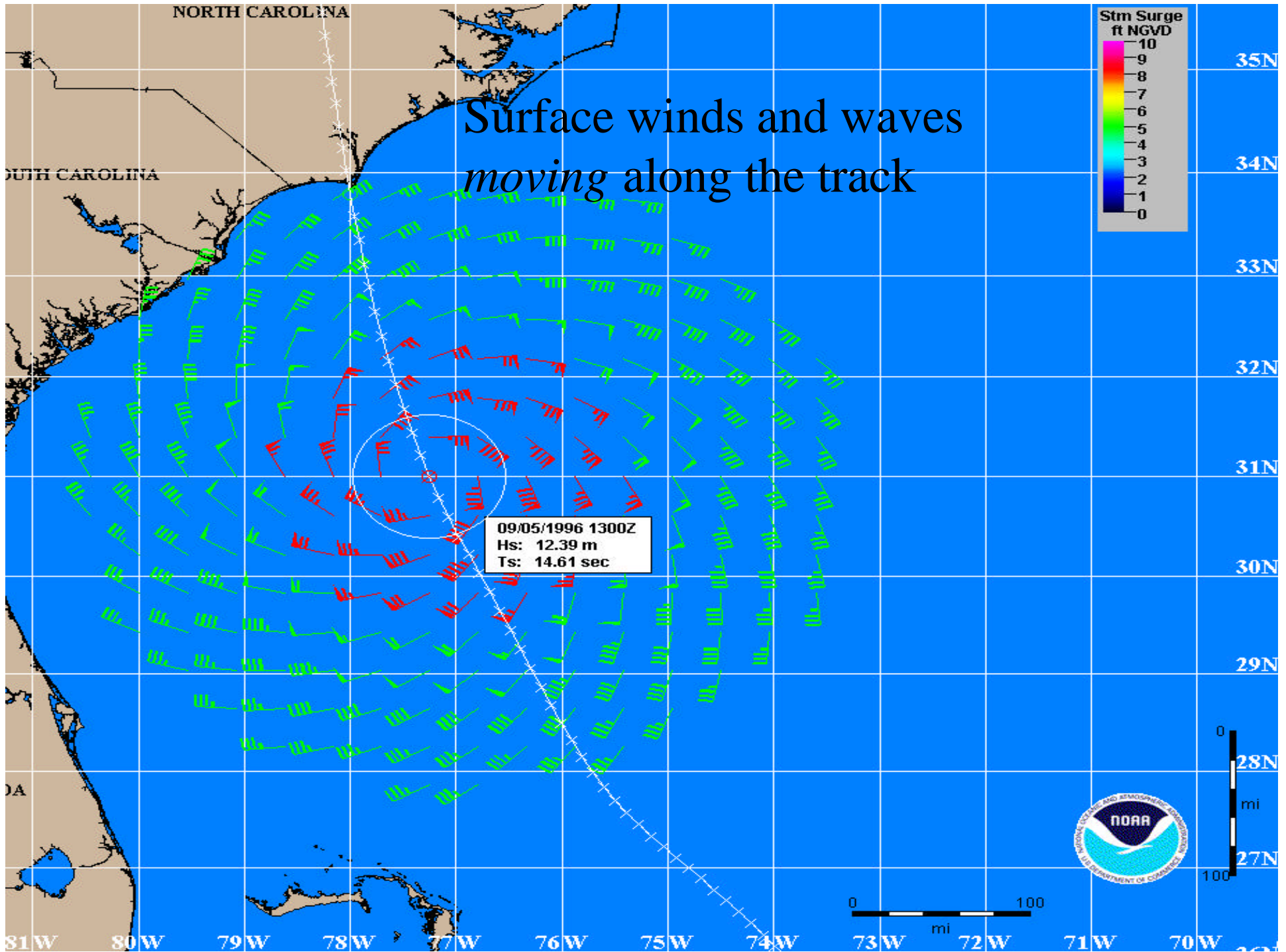
Winds,  
waves,  
water levels

September, 1999

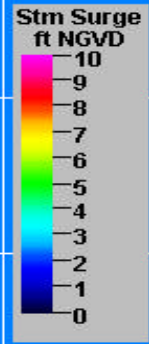
# GULF OF MEXICO SHORELINE PRELIMINARY STORM SURGE and WAVE EFFECT PROFILES FOR HURRICANE OPAL (1995)







# Surface winds and waves *moving along the track*



09/05/1996 1300Z  
Hs: 12.39 m  
Ts: 14.61 sec



## Meteorological Forcing of Ocean Surface waves

- A simple empirical hurricane wind-wave scheme based on JONSWAP fetch-limited wave growth  
I. Young (1988) proposed an equivalent fetch,  $F$ , such that

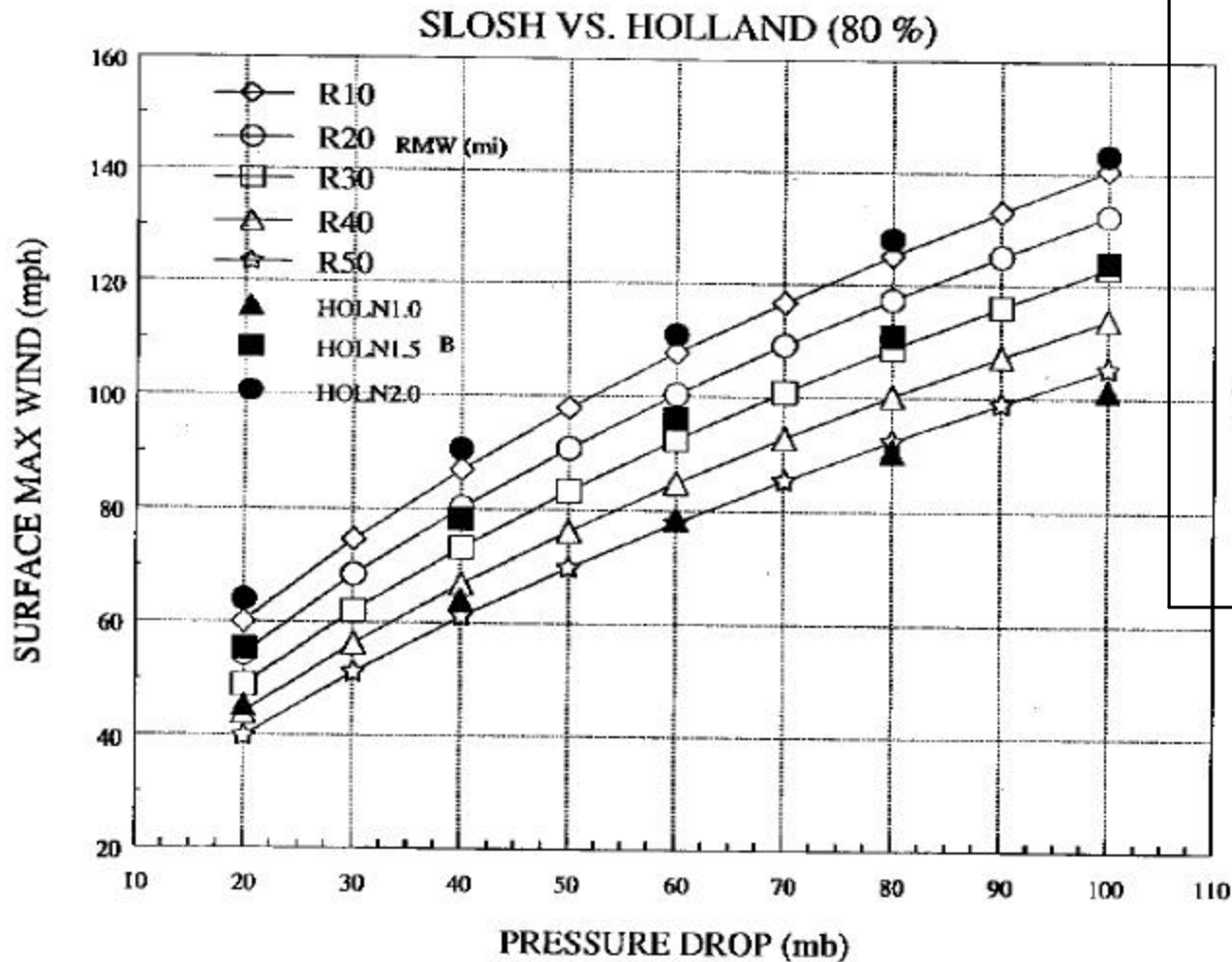
$$H_s = f(V_{\max}, F), \quad T_p \sim f(V_{\max}, F)$$

$F$  is tuned with wave model output and field data.

- . Among the sustained winds,  $V_{\max}$  is the key.

A Parametric cyclonic wind field

Surface wind vs. central pressure drop



A Stationary Storm  
 Myers & Malkin (1961)  
 Surface wind  $V$

$$\frac{1}{\rho_a} \frac{dp}{dr} = k_s \frac{V^2}{\sin\phi} - v \frac{dv}{dr}$$

and

$$\frac{1}{\rho_a r} \frac{dp}{d\phi} \cos\phi = fV + \frac{V^2}{r} \cos\phi - \frac{V^2}{\cos\phi} \frac{d\phi}{dr} + k_n V^2$$

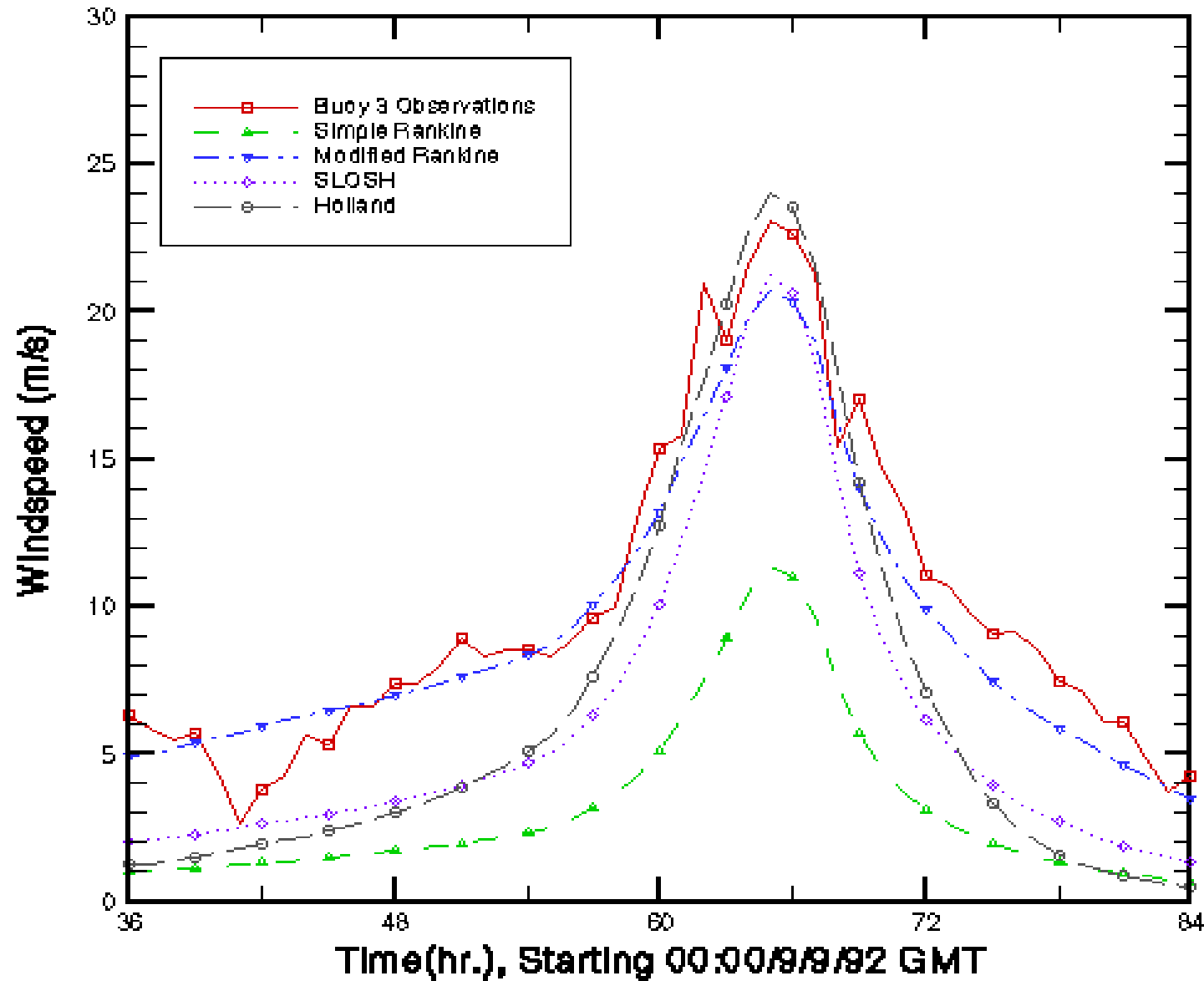
(1)

Solve  $p$  by assuming

$$V(r) = V_m \frac{2x}{1+x^2}$$

where  $x = r/R_{max}$ .

# Wind Model Comparisons vs. time







Hurricane wave based on equivalent fetch F (Young, 1988)

$$\frac{gH_s}{V_{\max}^2} = 0.0016 \left( \frac{gF}{V_{\max}^2} \right)^{0.5} \quad \text{For fetch-limited seas,}$$

$$\frac{F}{R^l} = aV_{\max}^2 + bV_{\max}V_f + cV_f^2 + dV_{\max} + eV_f + f$$

where  $R^l = 22.5 \times 10^3 \log R_{\max} - 70.8 \times 10^3$  (meters)  
(size of the dominant surface wind field)

Peak frequency  $f_p$ ,

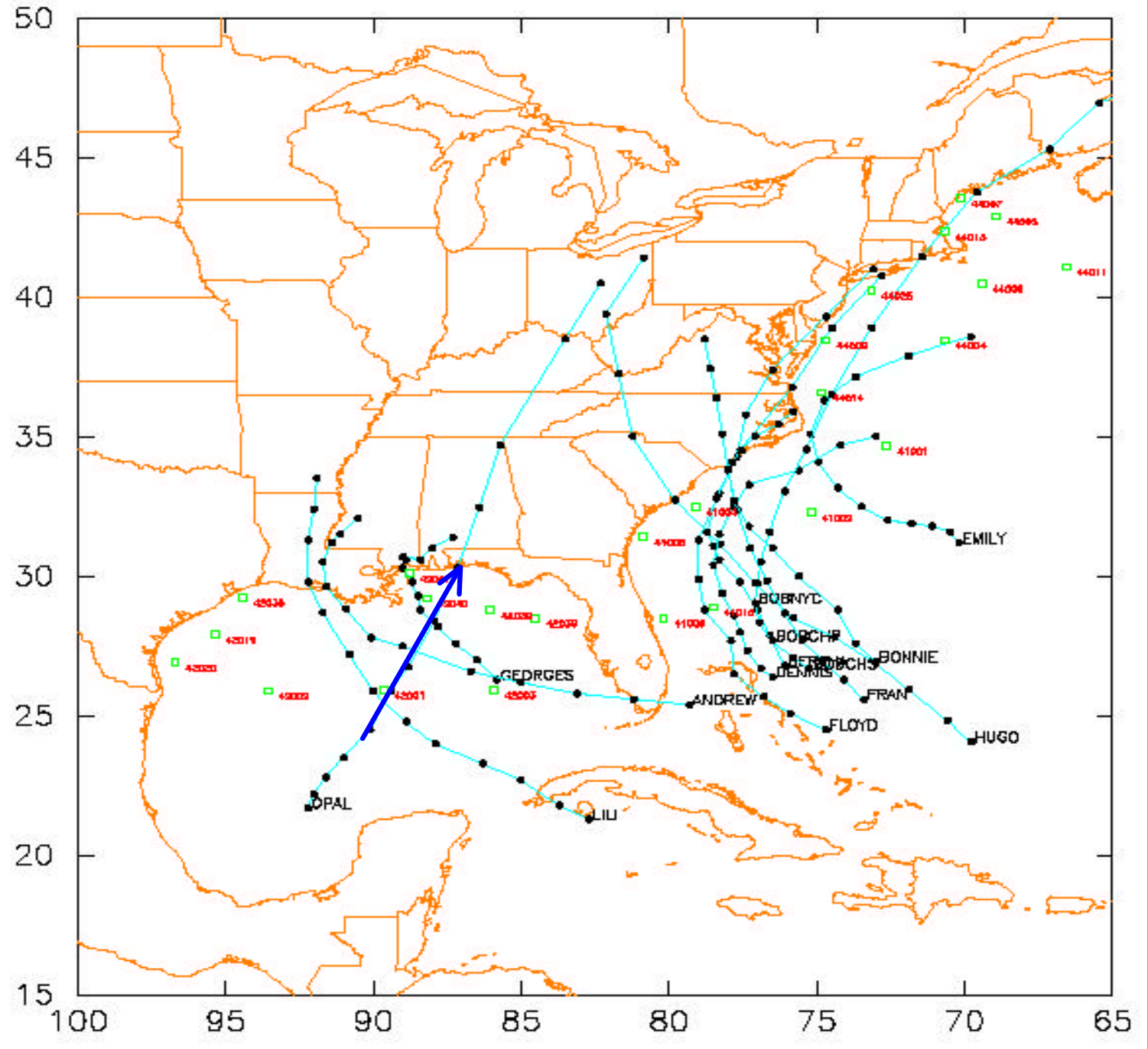
$$\frac{g}{2\pi f_p V_{\max}} = 0.045 \left( \frac{gF}{V_{\max}^2} \right)^{0.33}$$

LiLi (2002)

\*\*\* Tropical Cyclone WAVES and 12 ft seas radii  
RHS,LHS estimates the periphery of 12 ft seas  
RHS( the RHS of eye), LHS( the LHS of eye)

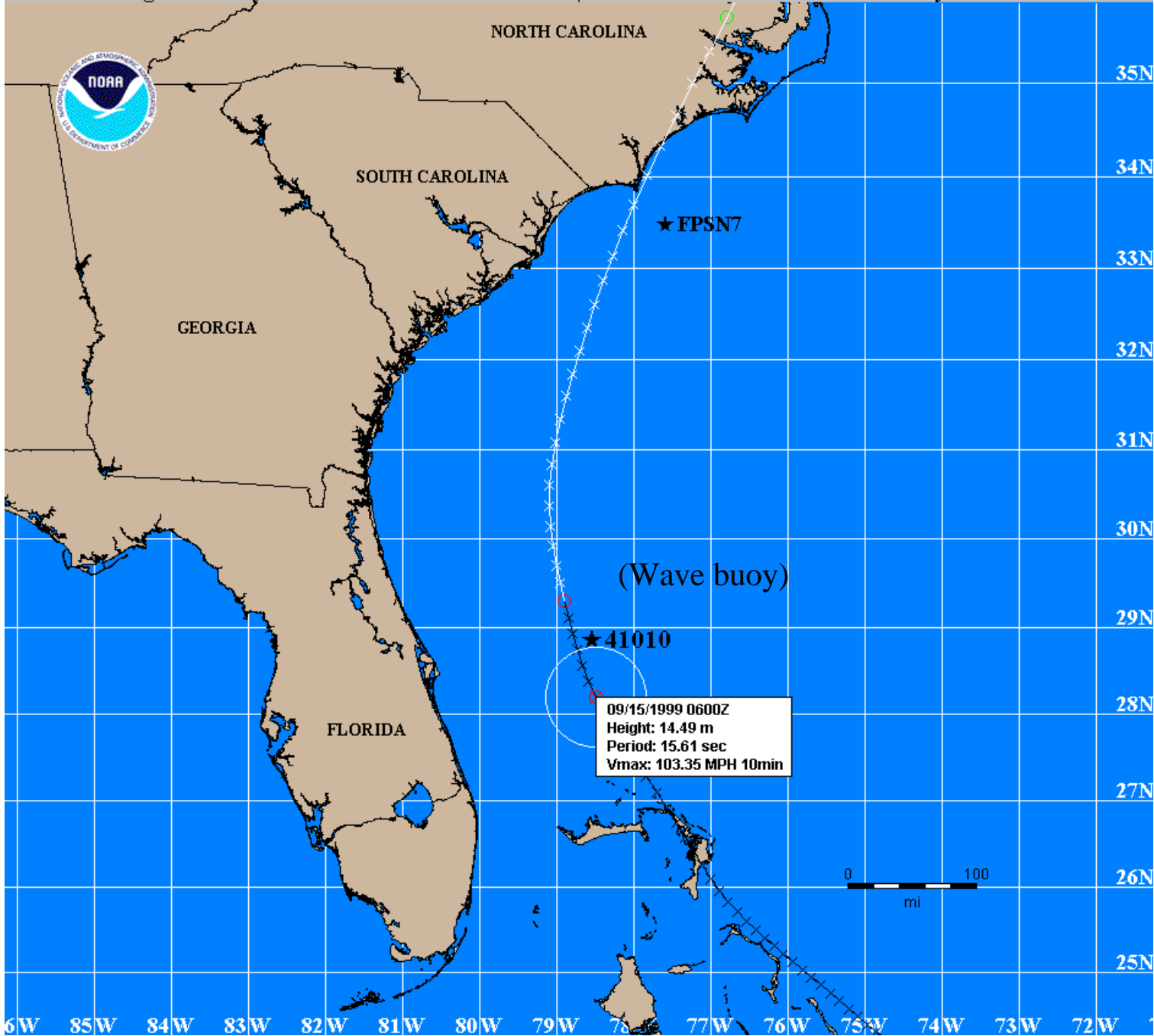
Maximum Waves estimated by calculating  $V_{max}$  ::

	PDROP	RMAX	VF	$V_{max}$	Wave-Ht	Wave-T	RHS	LHS
	(MB)	(NM)	(Knt)	(Mph)	(ft)	(SEC)	(NM)	
Seas	72.0	10.0	14.0	99.1	36.4	13.3	82.0	67.0



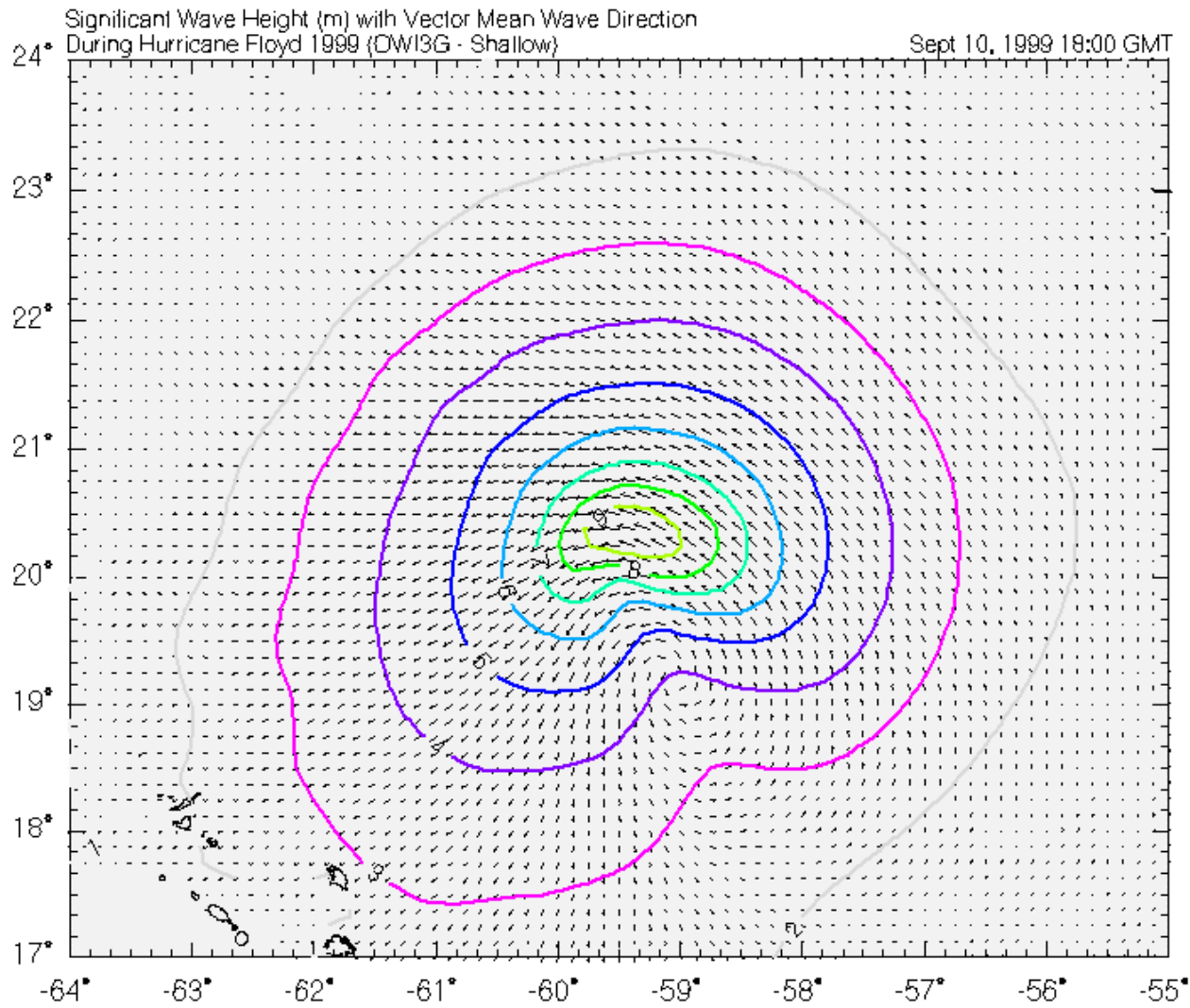
Basin: Wilmington NC / Myrtle Beach <ilm>

Storm: C:/slosh3.5/trkfiles/historic/floydilm.trk



Hurricane  
Floyd (1999)  
Choose time,  
yields Hs, T  
along the path





2-D wave  
using  
hourly  
winds  
 $dx=0.125$   
degree

2-D waves by E. Walsh (2001, JPO) on NOAA WP-3D. Using scanning radar altimeter.

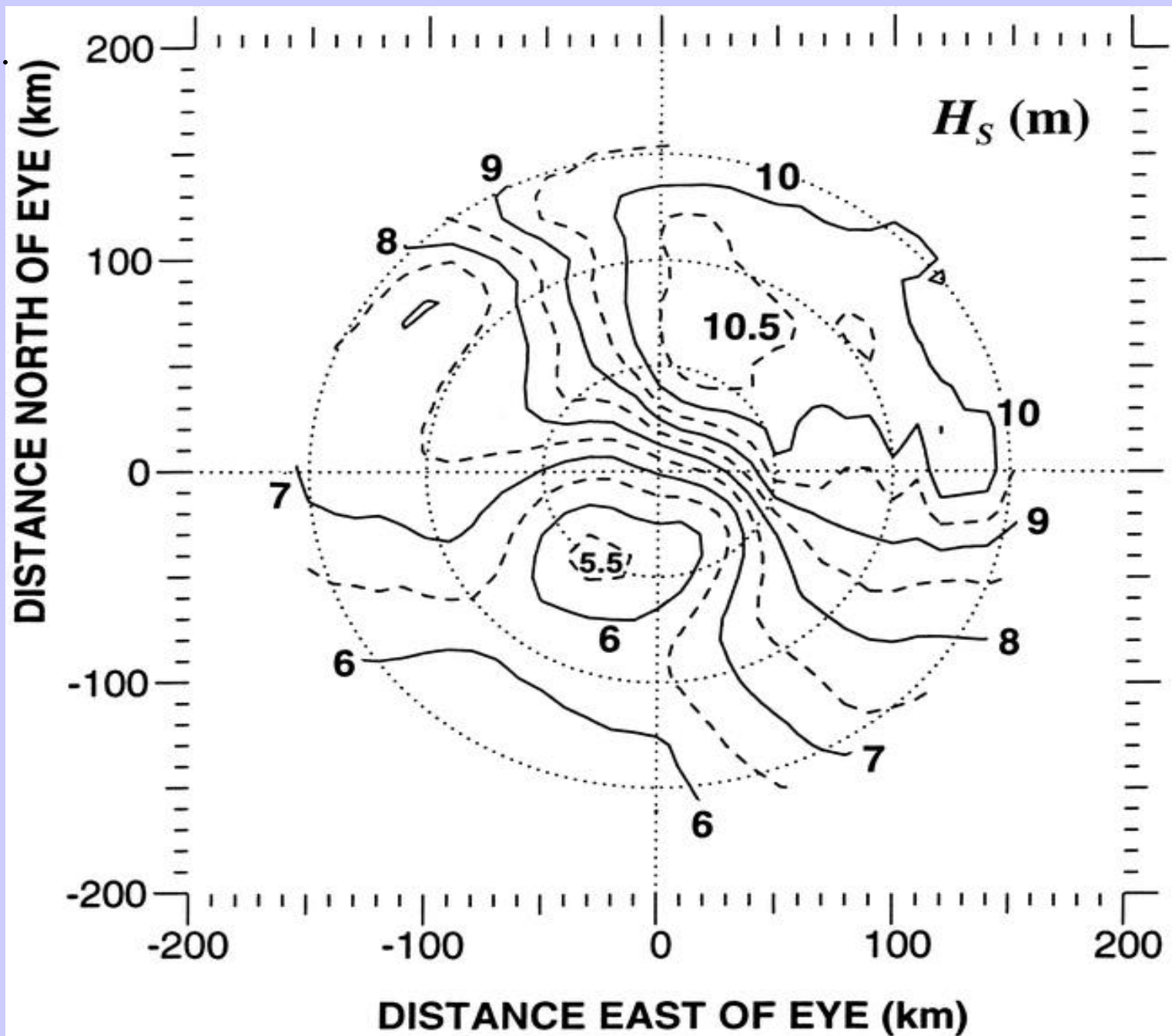
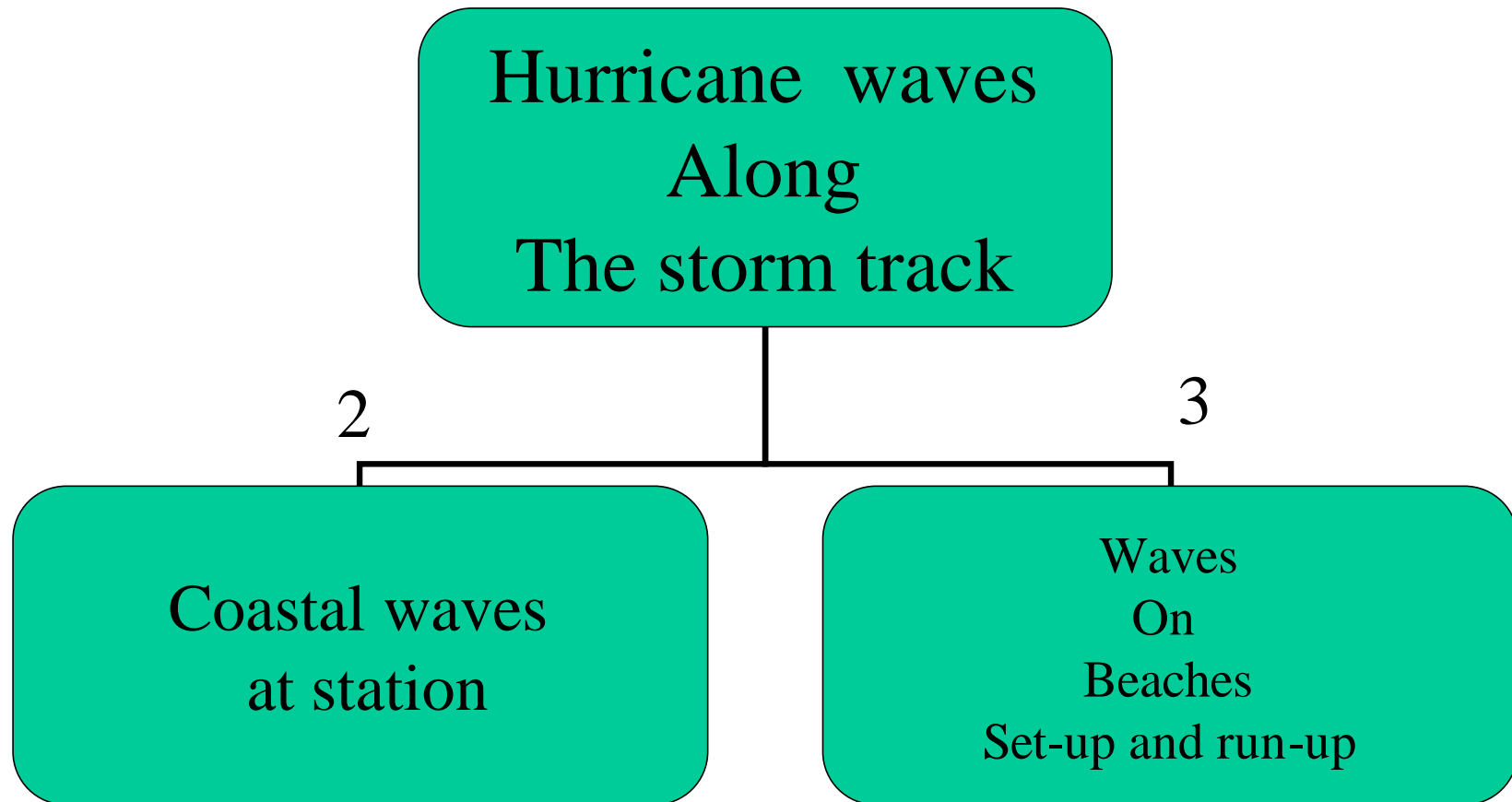


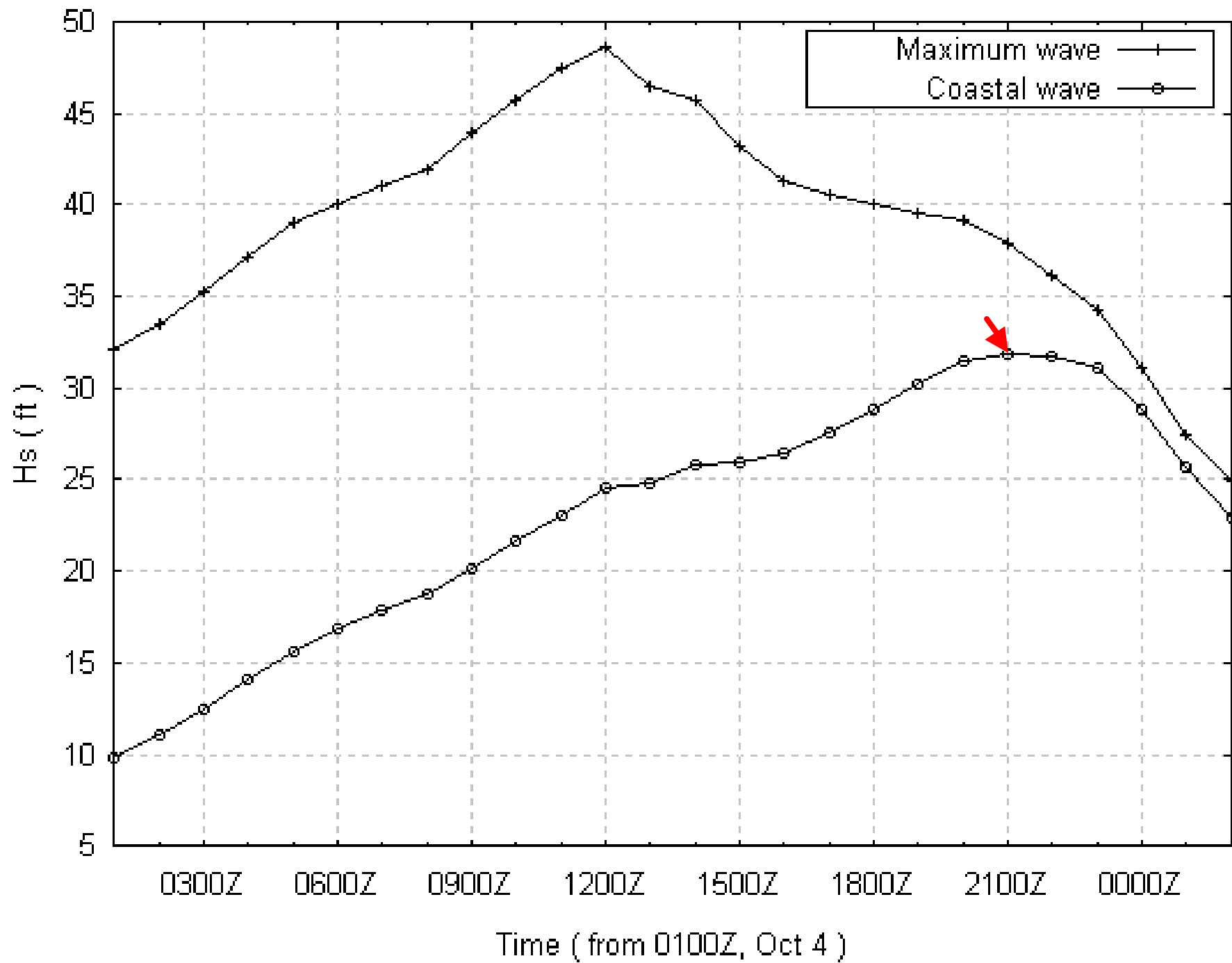
Table 1. Comparisons of wave heights( $H_s$ ) and wave periods( $T$ )

Hurricanes	Fran	Lili	Georges	Floyd	Bonnie	Iniki
track time	09/05/96 21Z	10/02/02 2030Z	09/27/98 16Z	09/15/99 08Z	08/24/98 21Z	09/11/92 18Z
Predicted $H_s$ (m)	11.50	11.12	10.48	14.24	10.46	4.95
Observed $H_s$ (m)	11.64	11.20	10.88	14.20	10.70	5.05
Predicted T (sec)	14.22	13.28	13.8	14.95	13.2	9
buoy <del>DMP</del> T(sec)	14.29	13.25	13.2	15.4	12.9	8.4
Data source	FPSN7	42001	42007	41010	Radar	51003

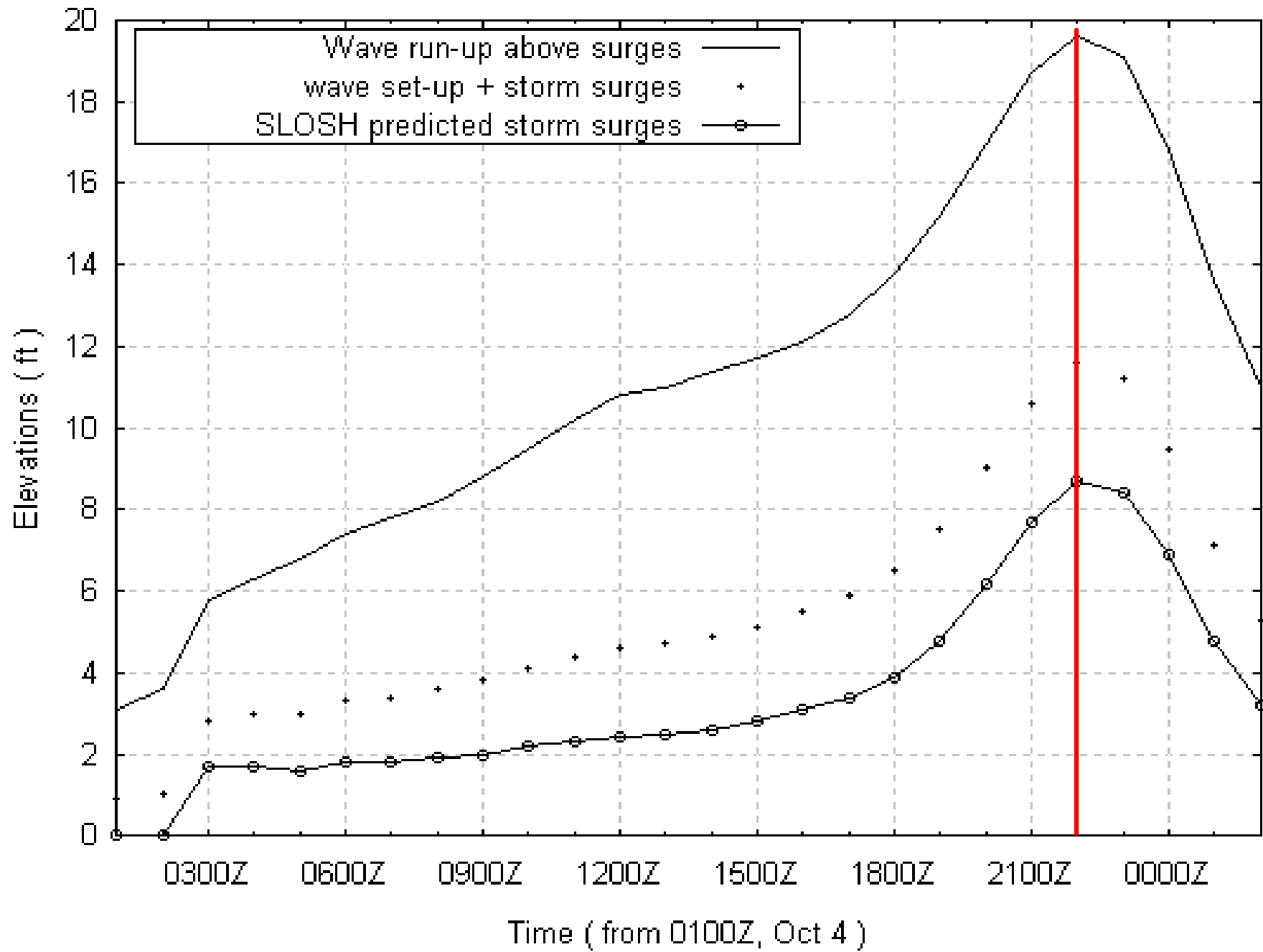
# Waves at Landfall site

## Computational procedures









## Remarks for operations:

Wave-surge must be included for slowly moving tropical cyclones making a landfall.

“A parametric hurricane wave model without high resolution grids can give accurate peak wave.”

- Wave induced surges ~ storm surge at the shore.
- Parametric model gives quick estimate of waves induced surge effects. The maximum surge occurs at the landfall, but waves are higher prior to landfall.
- The debris line is mostly attributed by wave run-up through the rise of sea water levels.

Waves on Beaches (Kweon and Goda, 1996):

$$\frac{\partial EC_{\xi}}{\partial x} = -\frac{K_d}{d_0} [EC_{\xi} - (EC_{\xi})_S] \quad (1)$$

Where E is the energy density and  $C_{\xi}$  the group velocity.

Using linear wave theory relationship, it leads to:

$$\frac{\partial [H^2 C_{\xi}]}{\partial x} = -\frac{K_d}{d} C_{\xi} [H^2 - H_s^2] \quad (2)$$

Where  $H_s = \Gamma d$ ,  $\Gamma = A^* (d/L)_o^{-1} [1 - \exp(-1.5 \pi (d/L)_o)]$ .

The wave-induced set-up  $\bar{\eta}$  is governed by time-averaged mean momentum equation:

$$\frac{d\bar{\eta}}{dx} = -\frac{1}{\rho g(d + \bar{\eta})} \frac{dS_{xx}}{dx}, \quad d = d_o + \bar{\eta} \quad (3)$$

And, wave run up follows Mase (1989, CEM)

$R = f(\text{beach slope and surf parameter})$